

Stratospheric Temperature Biases in the ERA5 reanalysis & plans for ERA6

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Biases in stratospheric temperatures in ERA5

- In '*model space*' - anomalies & analysis increments
- ERA 5.1 & the role of RO observations
- In '*observation space*' - mean first guess departures in ERA5, ERA-Interim & *proto-ERA6*

Southern winter polar bias in ERA5

- Exposed by anomalies & IRIS data

Exploiting information from RO observations back in time

- Weak constraint 4D-Var & model error forcing
- Using early sounding data (IRIS in 1970) to evaluate model error correction strategies

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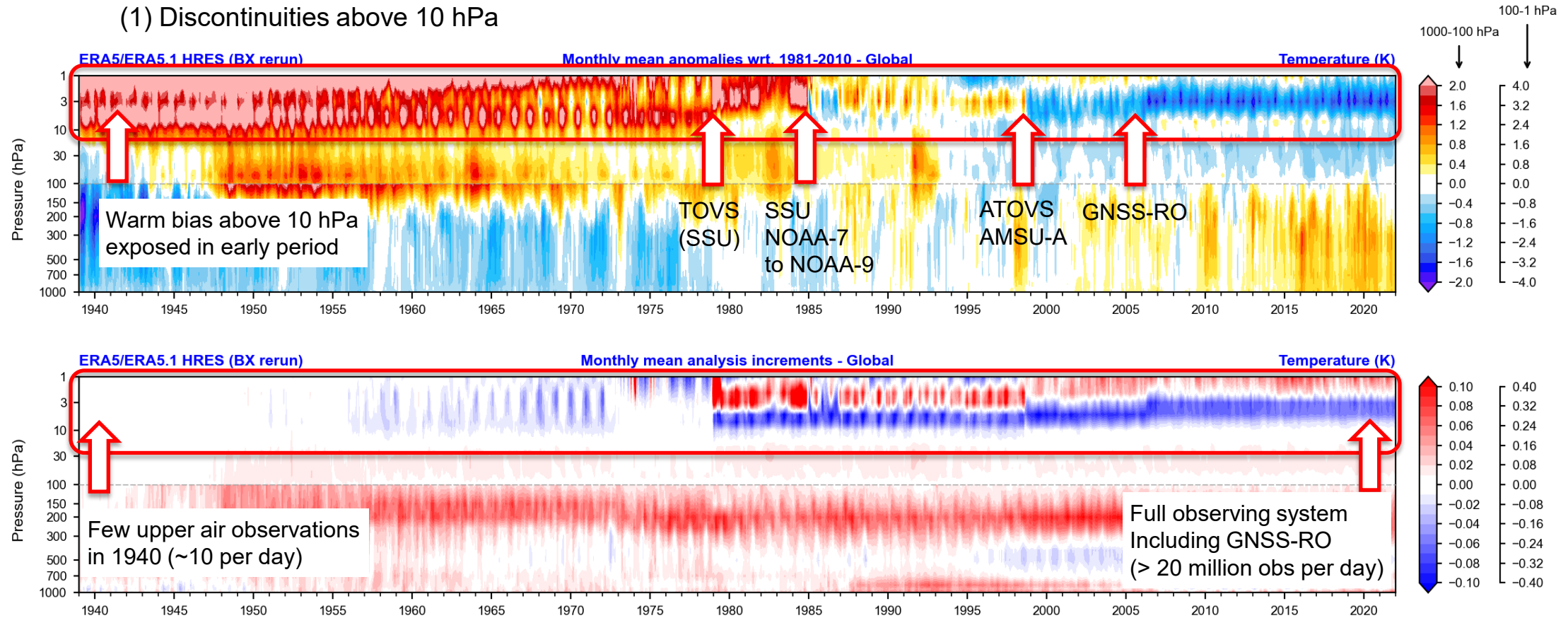
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Summary



Impacts of model and observation biases in ERA5

(1) Discontinuities above 10 hPa

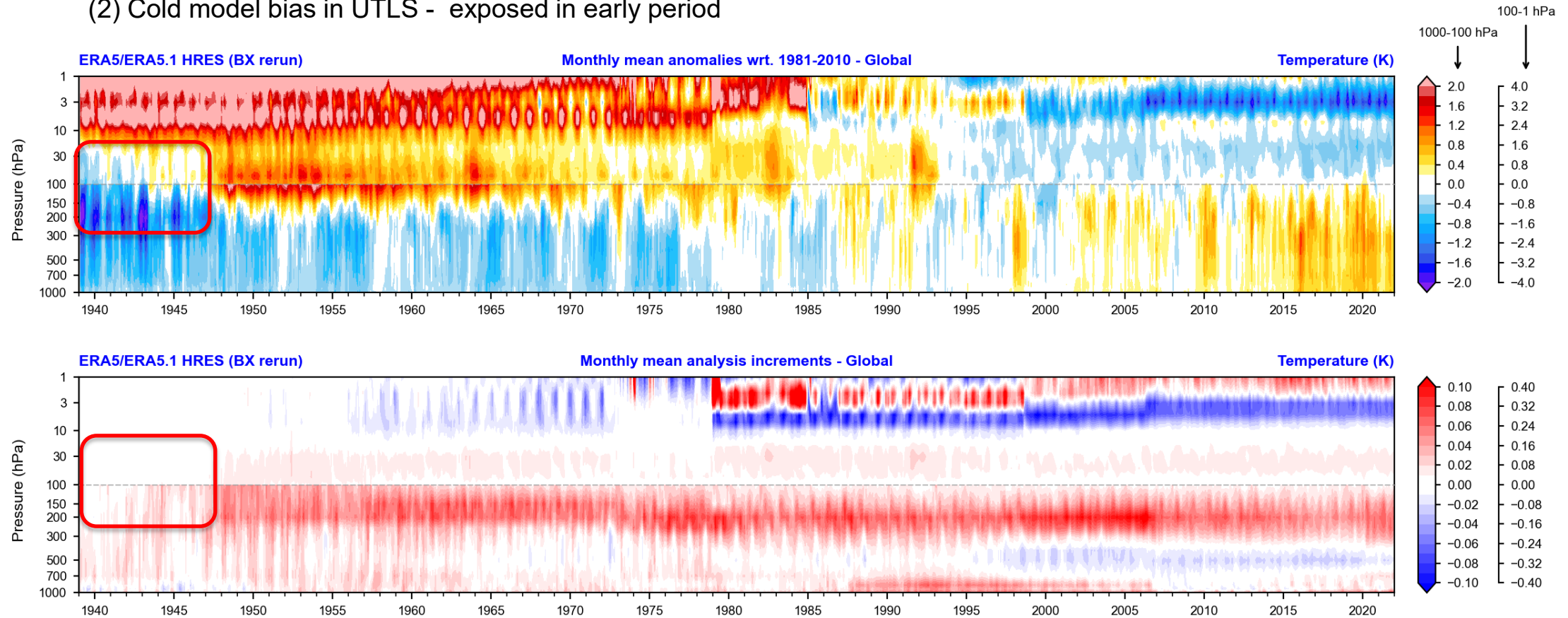


- General problems in reanalysis temperatures above 10 hPa well documented (see SPARC-RIP report 2021).



Impacts of model and observation biases in ERA5

(2) Cold model bias in UTLS - exposed in early period

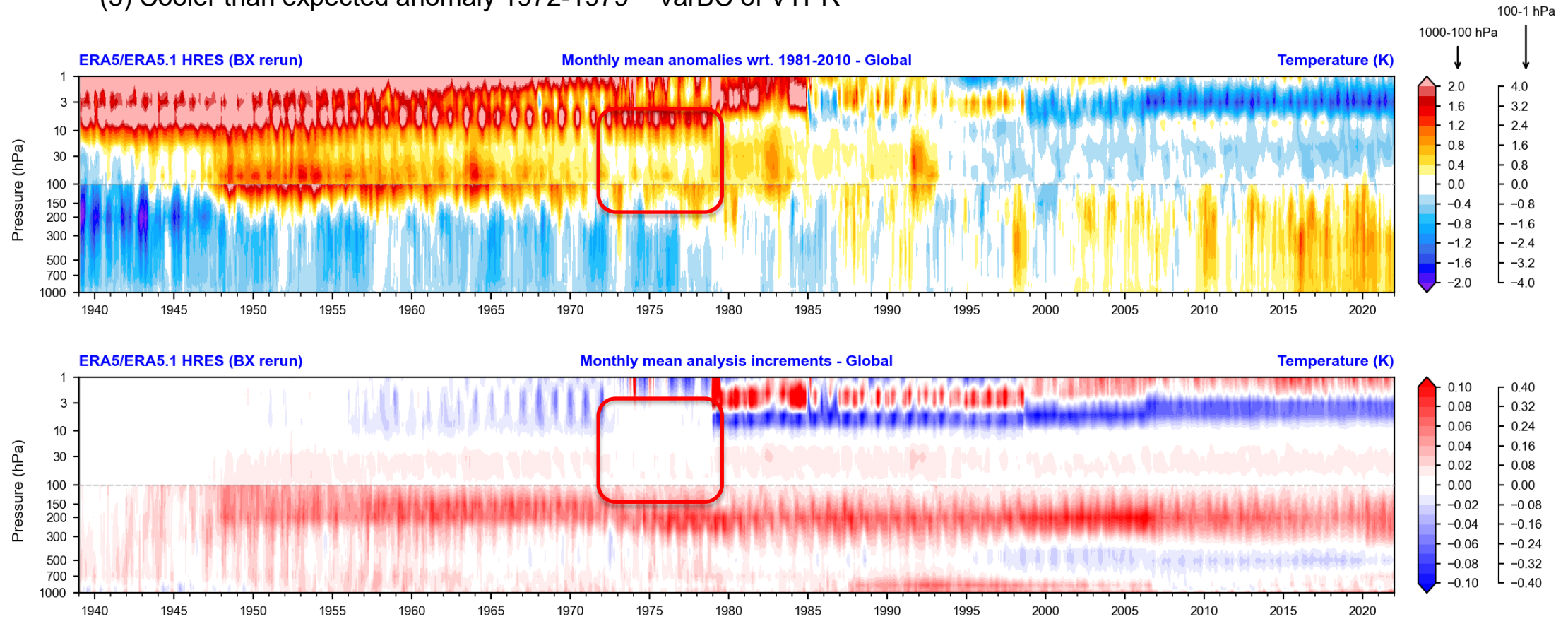


- Very few observational constraints on stratospheric temperature analysis in the early 1940s – so UTLS cold bias is exposed.
- Analysis increments in 10-200 hPa layer very small 1940 (< 20mK above 100 hPa as a global mean)



Impacts of model and observation biases in ERA5

(3) Cooler than expected anomaly 1972-1979 – VarBC of VTPR



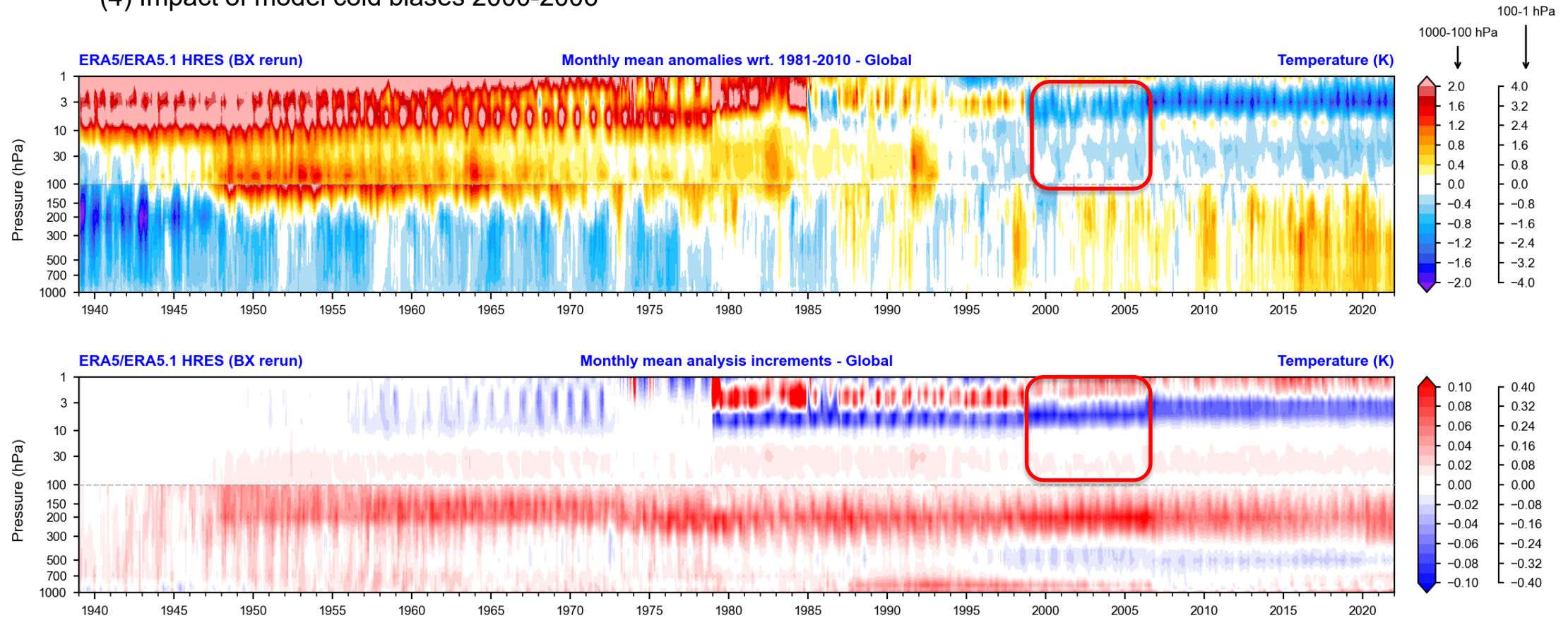
- General problem foreseen & analysed in Eyre (QJ, 2017): with VarBC, if radiances are dominant (cf anchors) model bias is reinforced
- VTPR channels 1 & 2 bias corrected using VarBC – reinforcing model cold bias
- Despite clear benefits (from assimilating VTPR) in improving synoptic analysis (earlier slide) – mean state exhibits a discontinuity.
- VTPR exhibits significant radiometric and spectral errors ⇒ we need VarBC !



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Impacts of model and observation biases in ERA5

(4) Impact of model cold biases 2000-2006



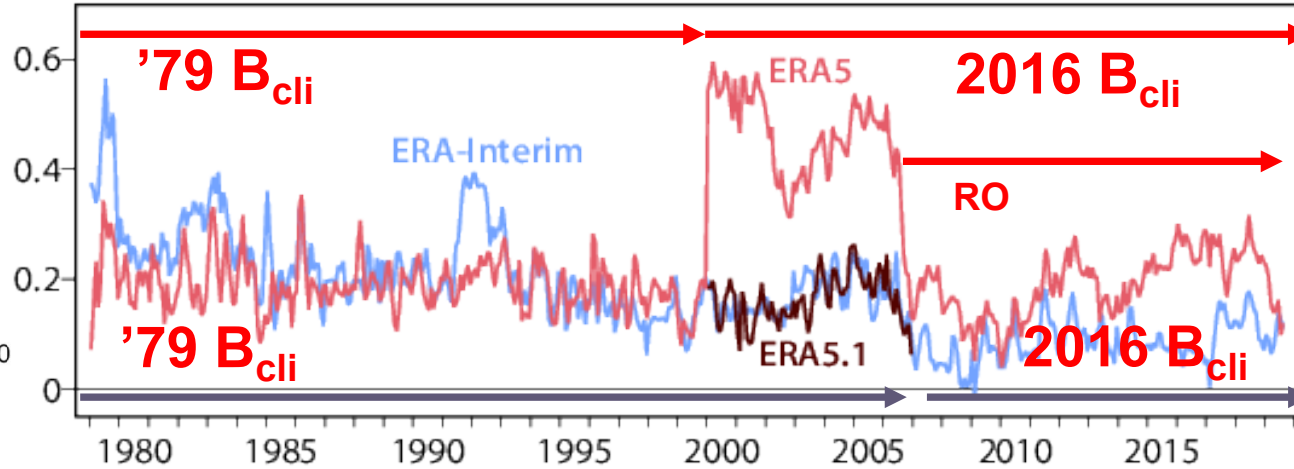
- ERA5 and ERA5.1: See next slide



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The improved mean state for stratospheric temperature in ERA5.1

Global-mean ob-bg for 60-40hPa radiosonde temperatures (K)



ORIGINAL ERA5 PRODUCTION

ERA5.1 PRODUCTION

Monthly average observation-background differences from 1979 onwards for all assimilated bias-adjusted radiosonde temperature data (K) between 40 and 60 hPa, for ERA-Interim, ERA5 (based on 1979-B_{cli} before 2000 and 41r2-B_{cli} afterwards) and ERA5.1 (using 1979-B_{cli} from 2000-2006).

Hersbach, H. et al., 2020, [doi:10.1002/qj.3803](https://doi.org/10.1002/qj.3803)

- ERA5.1 provides an improved mean state for stratospheric temperature.
- In the troposphere the difference between ERA5 and ERA5.1 is very small.

(see A. Simmons *et al*, ECMWF Tech Memo 859, Jan 2020)

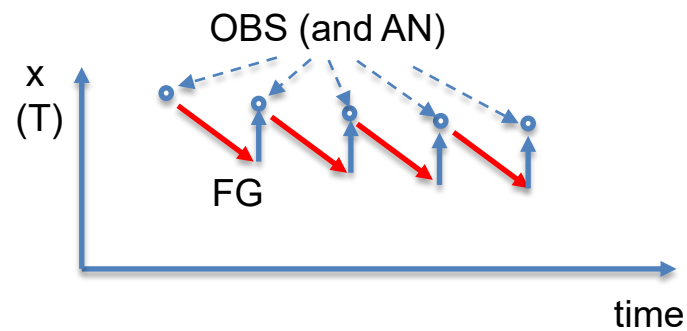
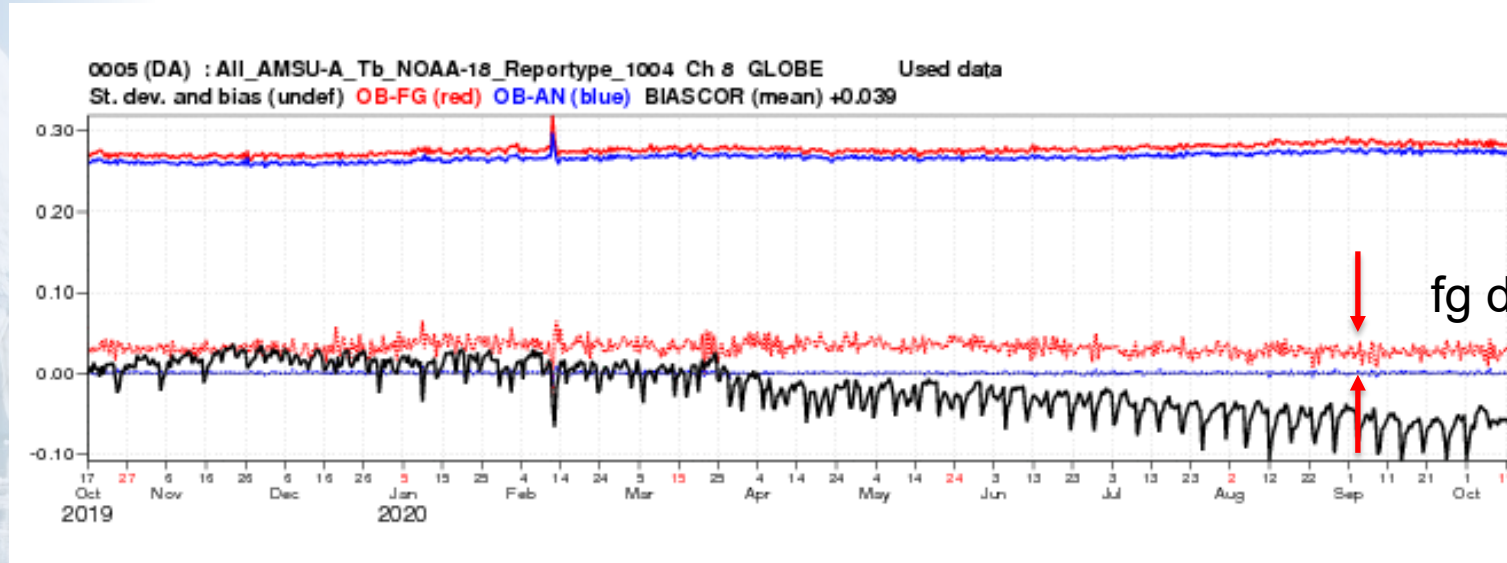




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Model error manifested in biased first guess departures

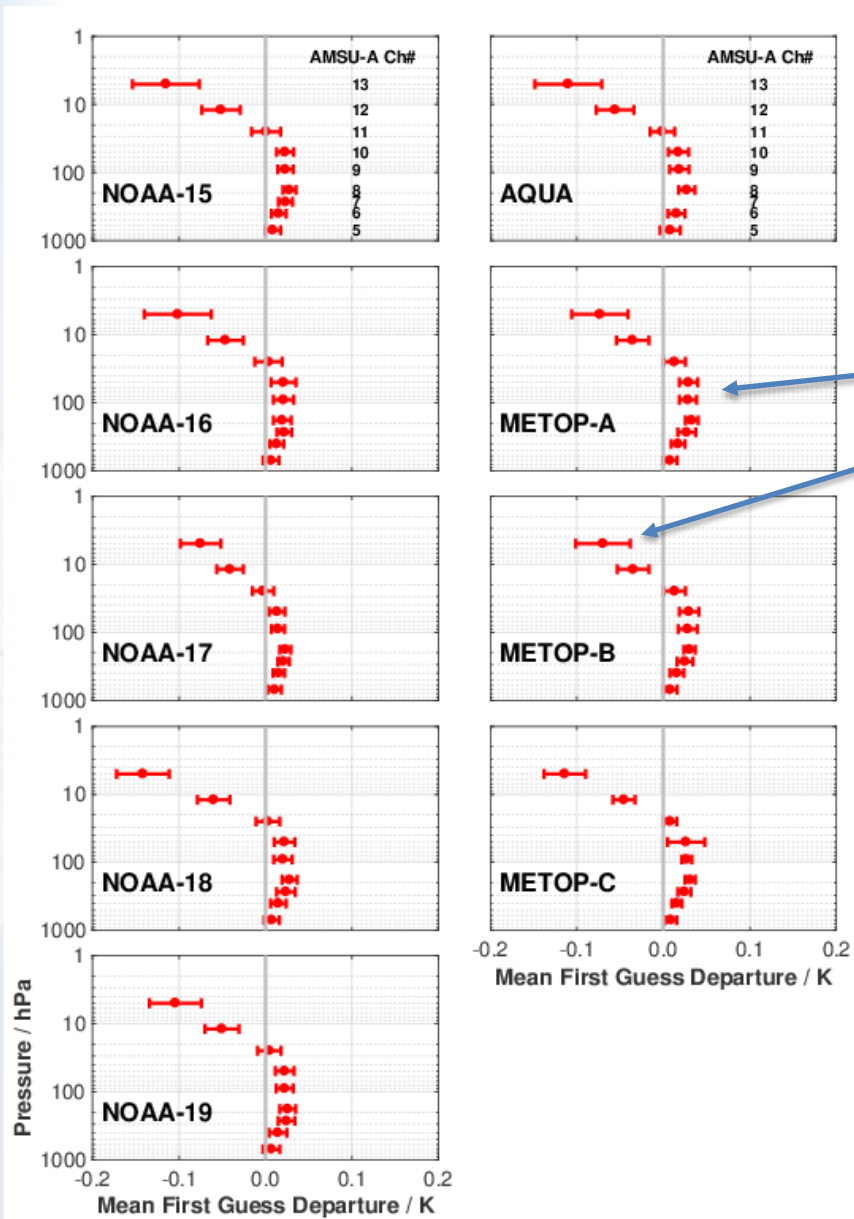
NOAA-18 AMSU-A8





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Model Error / AMSU-A Mean first guess departures in ERA5



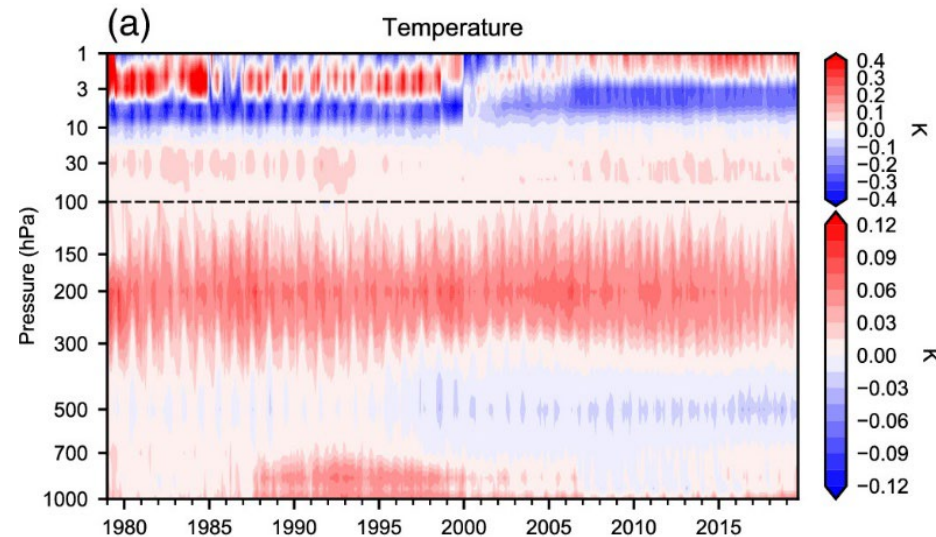
ERA5 mean first guess departures shown for AMSU-A

Error bars represent ($\pm 1\sigma$) spread over the lifetime of each sensor

Consistent picture of :

- a cold model bias mid-trop to mid-strat
- a (larger) warm model bias above 10 hPa

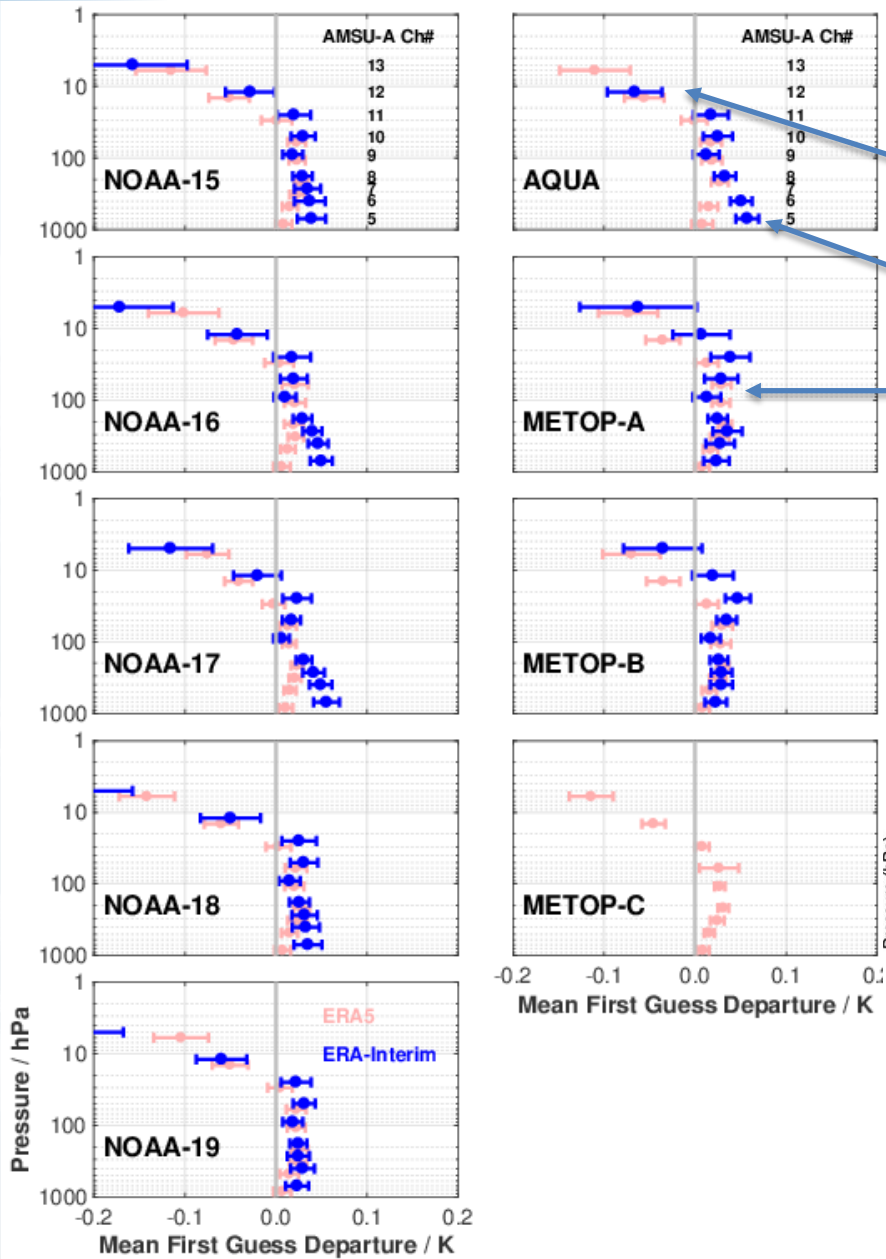
Broadly consistent with analysis increments in ERA5
(below, from Fig 16, Hersbach et al, 2020)





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Model Error / AMSU-A Mean first guess departures in ERA-Interim



Indications that ERA-Interim:

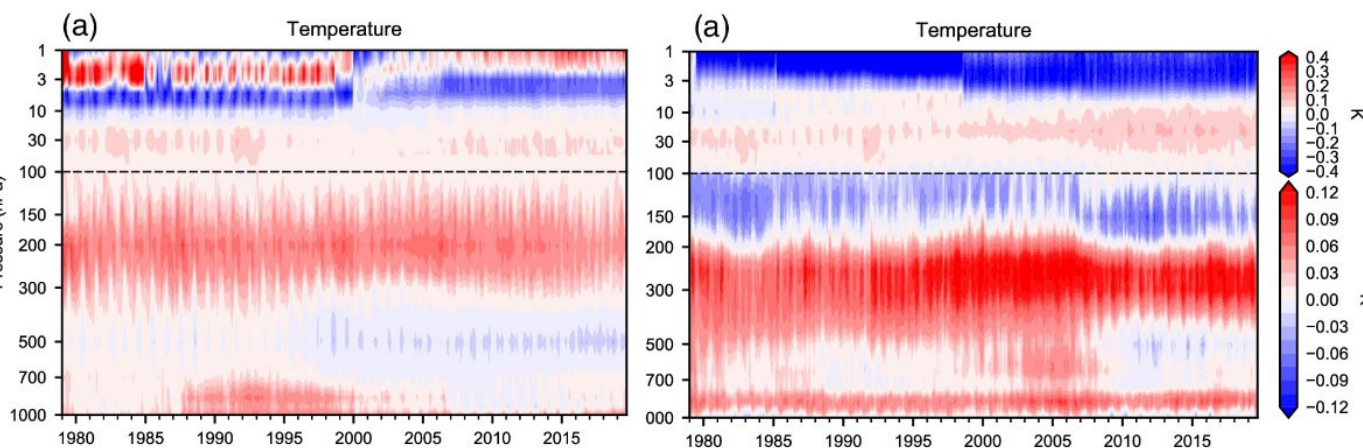
Exhibits similar biases (to ERA5) above 10 hPa

Exhibits larger biases below 200 hPa

Exhibits smaller biases around 100 hPa.

ERA5
analysis increments

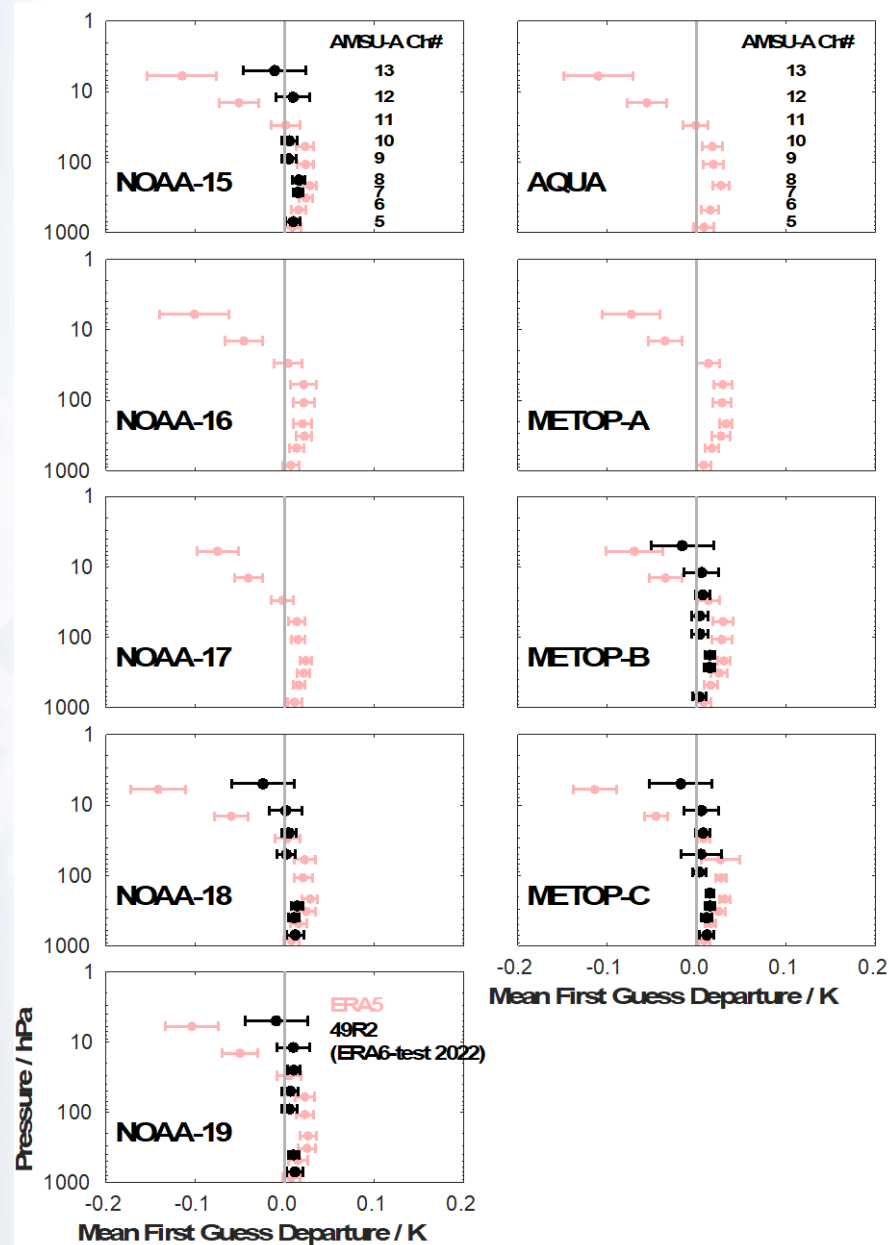
ERA-Interim
analysis increments





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Model Error / AMSU-A Mean first guess departures in proto-ERA6 testing (CY49R2)



- Several improvements in analysis of the stratosphere since 2016:
 - Weak constraint 4D-Var
 - Improvement in dynamics
- Statistics shown based on JJA 2022 49R2 experiment
- Tco639 (18 km resolution) - ERA6 production will be 14 km
- Overall – forecast model in better agreement with observations



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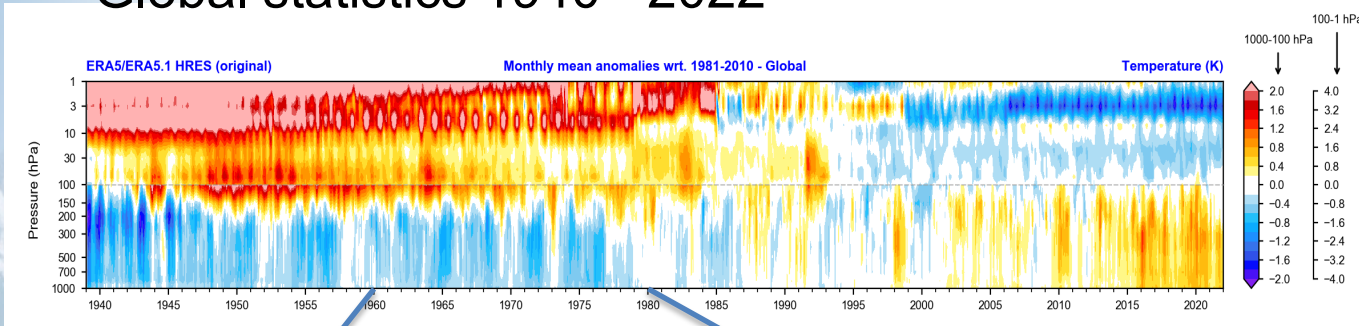
Summary / Future Perspectives



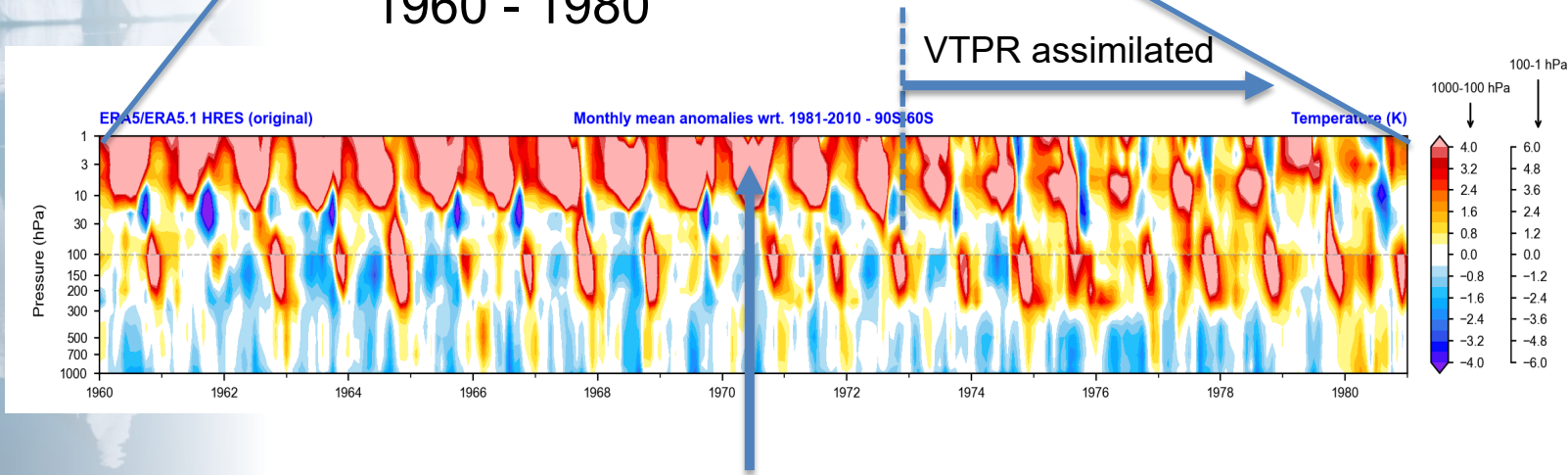
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Upper stratospheric biases in ERA5: Temperature anomalies relative to ERA5 climate

Global statistics 1940 - 2022



Southern polar statistics 1960 - 1980

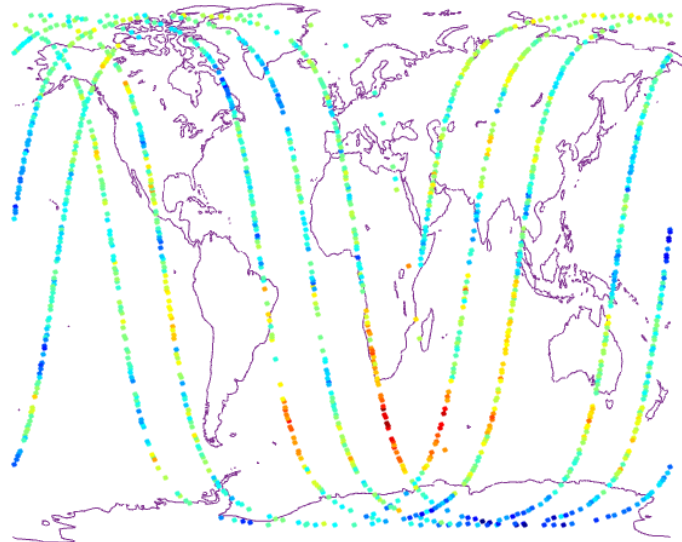


IRIS experiments

- Generally, ERA5 temperature analyses above 10 hPa exhibit biases and discontinuities
- Particularly large biases evident in southern polar winter ($\gg 6K$ in the plot shown)
- Repeatable from year-to-year (before 1972)
- Reduced following the assimilation of VTPR data (Nov 1972 - Jan 1979)



Initial experiments assimilating IRIS in the IFS



Typical 12 hour coverage

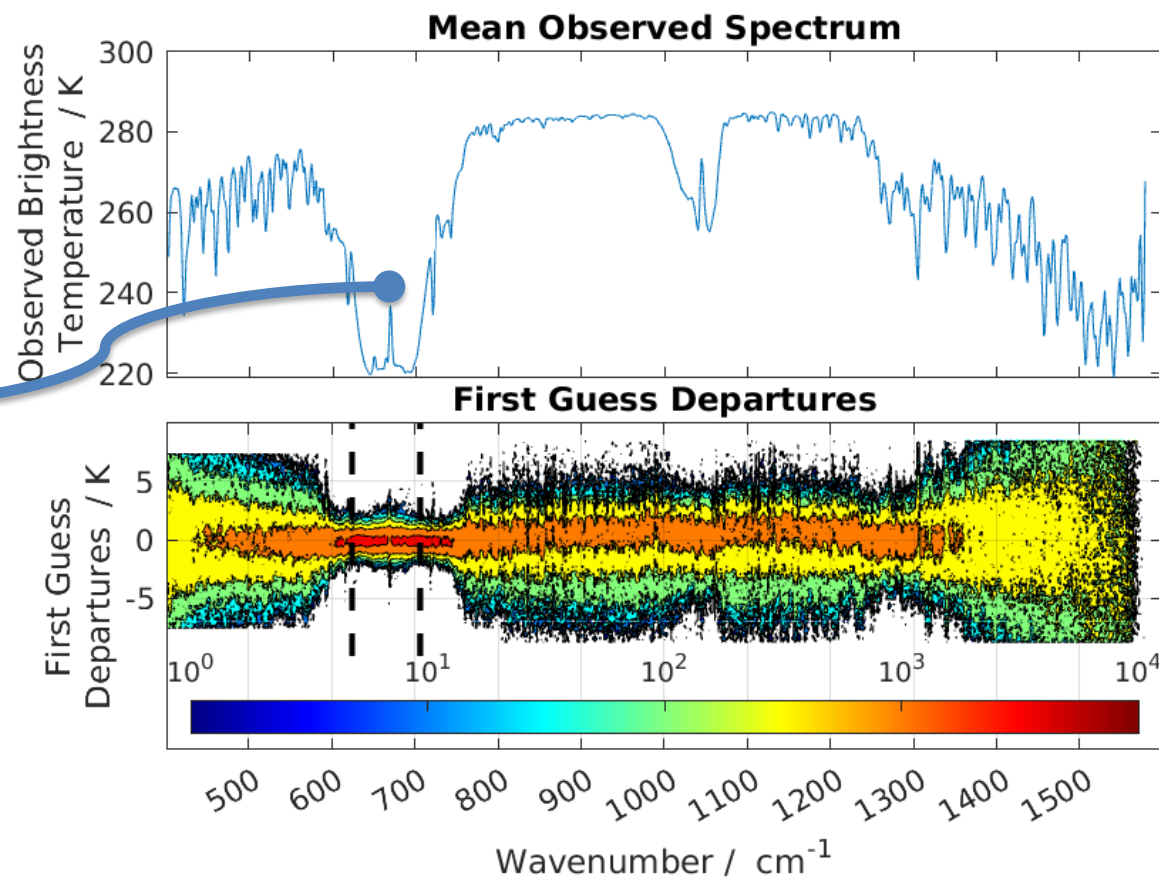
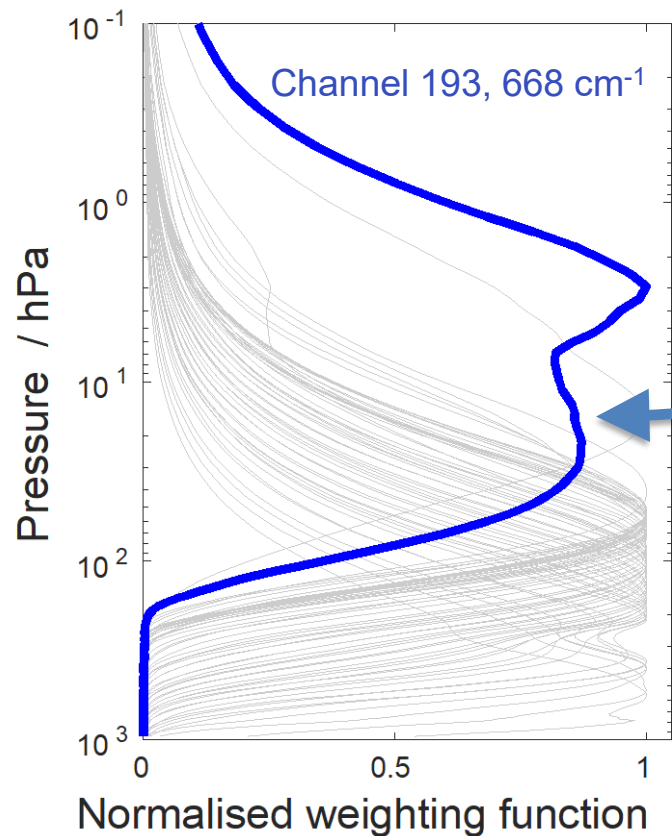
- Operated on Nimbus-4, from April 1970 – January 1971
- Nadir only observations. Spectral range 400 - 1600 cm^{-1}
- Resolution: 2.53 cm^{-1} to 2.69 cm^{-1}
- 94 km footprint
- 13 s measurement time
- Coverage to 80°N to 80°S (rely on **B** to propagate information to poles)

Daily time coverage / %

Year	Month	01	02	03	04	05	06	07	08	09	10	11	12	13	14	15	16	17	18	19	20	21	22	23	24	25	26	27	28	29	30	31
1970	04									18	63	0	0	2	95	88	77	86	2	79	93	88	85	85	5	83	0	1	86	79	76	
	05	82	3	85	5	92	79	85	85	93	93	93	93	79	52	84	93	93	5	93	86	81	0	92	86	93	90	90	82	79	5	88
	06	90	0	90	77	59	59	71	70	86	63	77	79	93	93	91	93	6	90	0	3	93	93	93	93	5	93	91	0	91	93	
	07	93	1	90	93	93	85	85	5	0	89	89	68	92	73	4	76	94	3	0	0	4	86	5	94	94	79	79	1	0	0	89
	08	85	79	88	1	0	76	9	91	85	86	82	90	86	94	80	0	55	5	94	2	59	2	79	80	86	5	0	0	2	82	47
	09	86	88	2	88	95	80	95	86	95	2	0	91	89	80	94	5	75	0	47	36	3	0	0	0	96	94	88	89	93	94	
	10	2	85	61	63	58	85	95	86	6	0	0	0	4	21	93	80	5	89	1	90	63	3	88	94	88	4	4	85	78	0	0
	11	93	91	91	85	0	0	60	86	83	0	93	81	93	97	1	91	92	94	89	93	98	93	82	54	2	96	93	69	82	81	
	12	71	70	68	1	69	77	76	0	77	0	1	63	66	61	0	61	70	69	0	76	72	72	73	68	74	3	71	72	69	69	75
	1971	01	0	70	64	73	63	69	0	0	53	0	0	0	0	0	0	0	46	52	0	36	42	46	0	35	44	46	0	40	40	48



Investigating biases using early hyperspectral sounding data (Nimbus-4 IRIS, 1970)



- IRIS data has been shown to be valuable in improving SH analysis quality (April 1970 – January 1971)
- Valuable for assessing biases in ERA5 in previously unobserved regions (eg S. Polar upper stratosphere)
- Highest peaking channel is particularly valuable



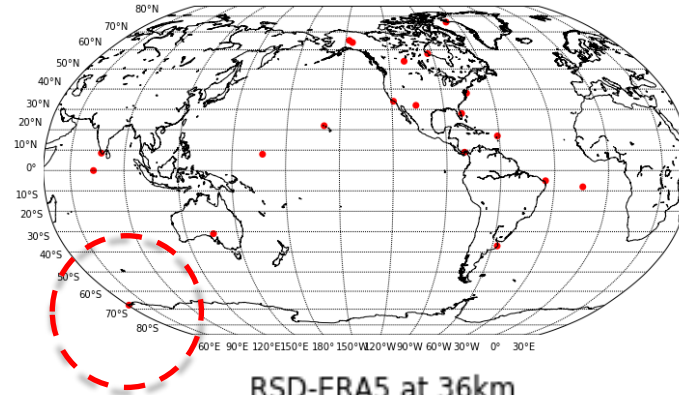
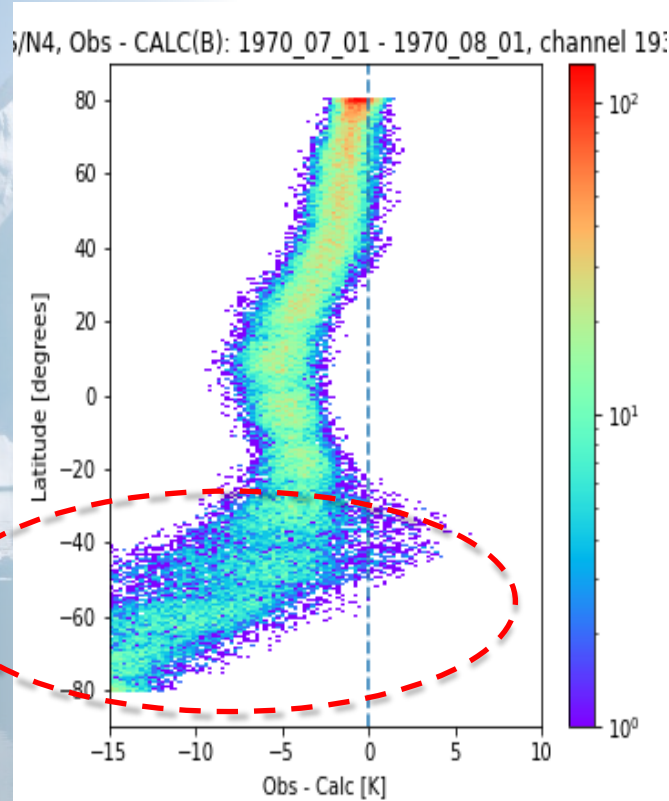
Evaluation of IRIS radiances relative to ERA5, rocketsondes & SIRS radiances (Andrzej Klonecki *et al* , Spascia)

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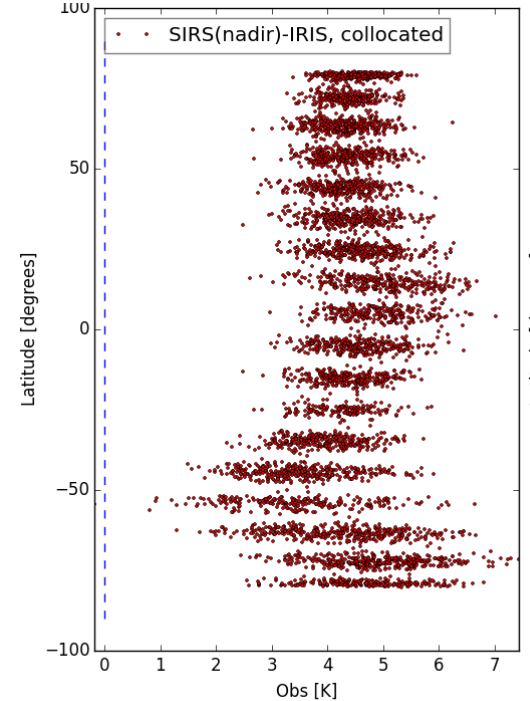
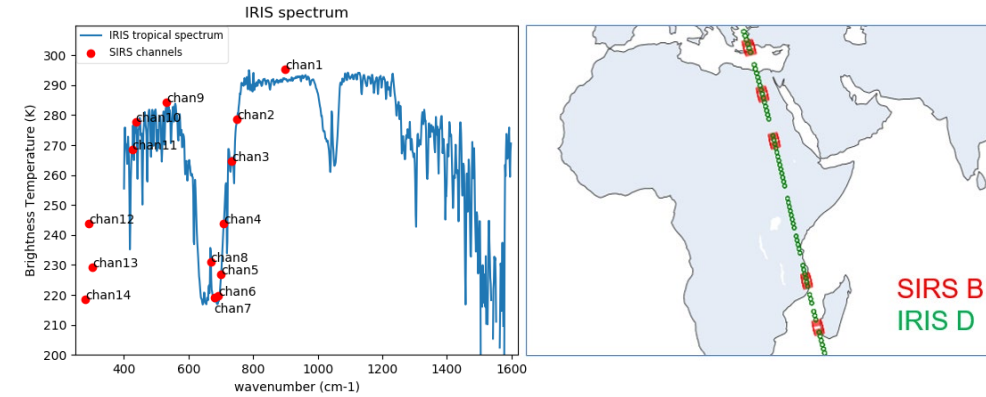
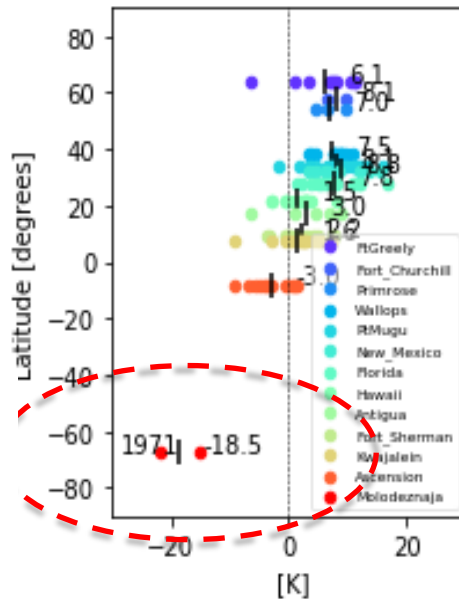
IRIS Obs vs ERA5 channel 193 peaking at 3 - 40hPa

rocket sondes vs ERA5

Nimbus-4 IRIS vs SIRS



RSD-ERA5 at 36km
1970-07-01-1970-07-31



SIRS / IRIS
(ch 8 / ch 193)
colocations for
July 1970



In summary, relative to :

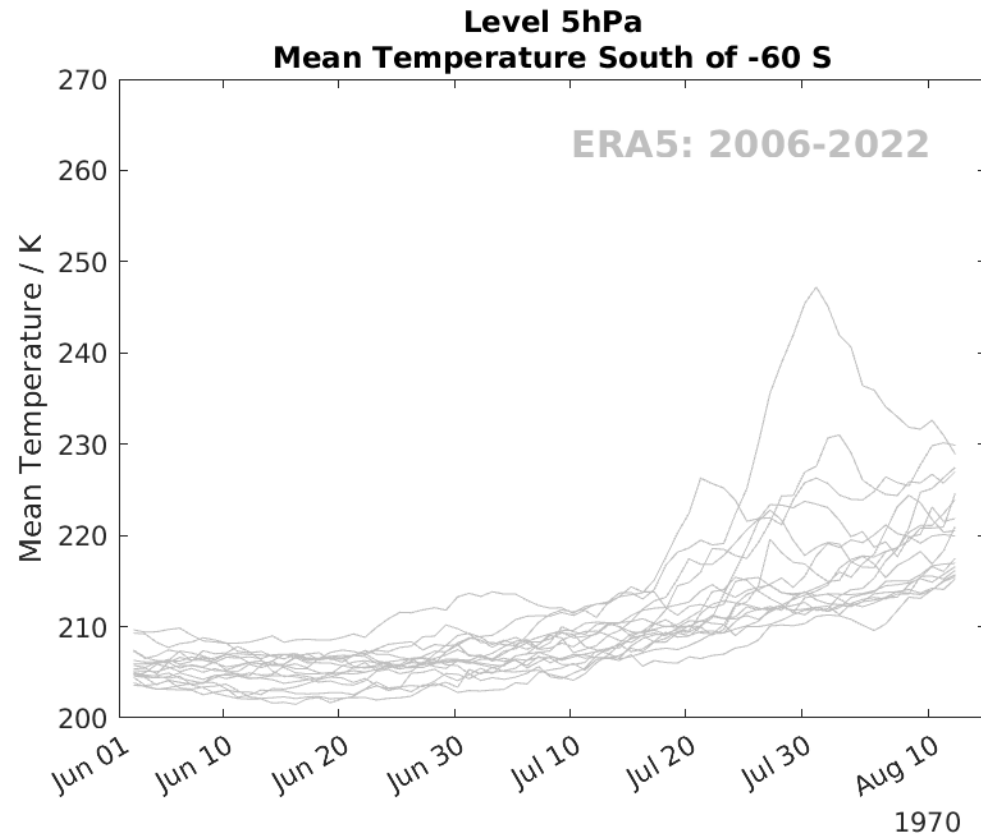
- Its own climate (1981-2010);
- IRIS observations;
- SIRS observations; and
- Rocket-sonde data

ERA5 exhibits a warm bias, at 36km / 5hPa, of ~15K



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Impact of assimilating IRIS on southern polar stratospheric biases

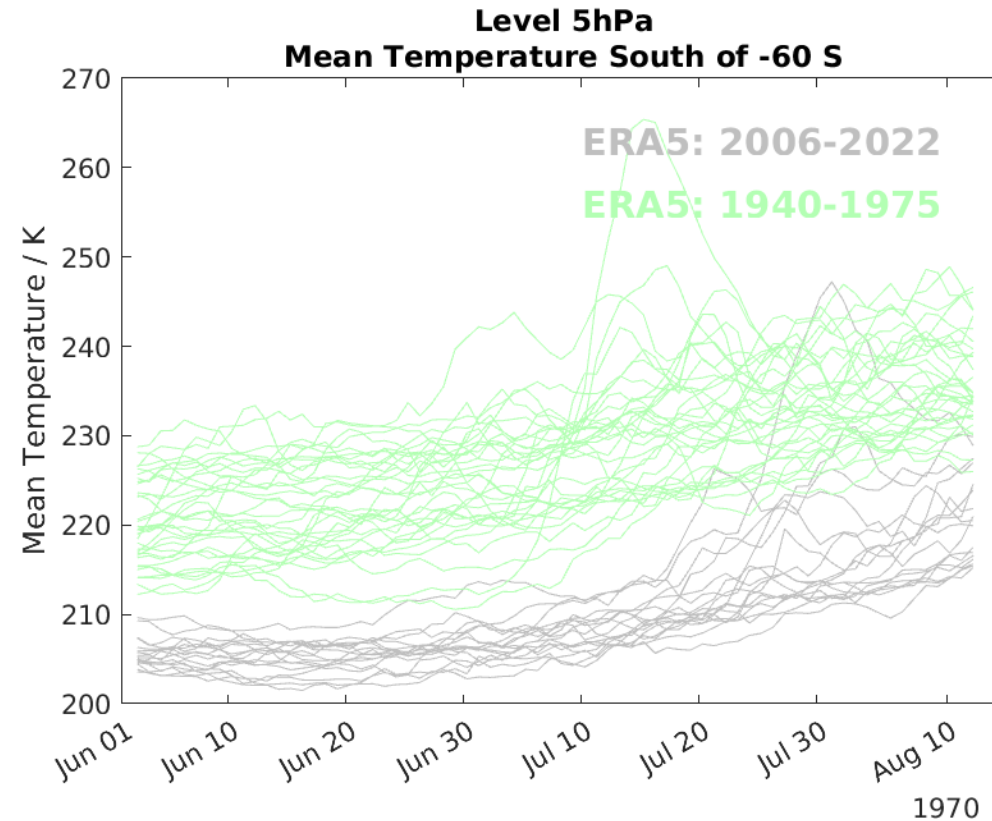


- During the GNSS-RO era (2006 -) the stratospheric temperature analysis is realistic



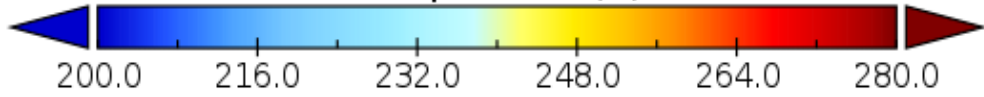
Climate
Change

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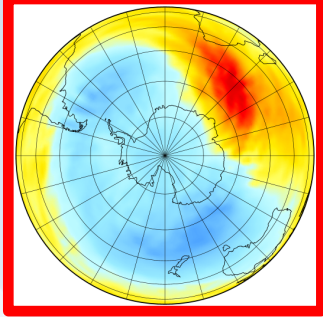
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Temperature (K)

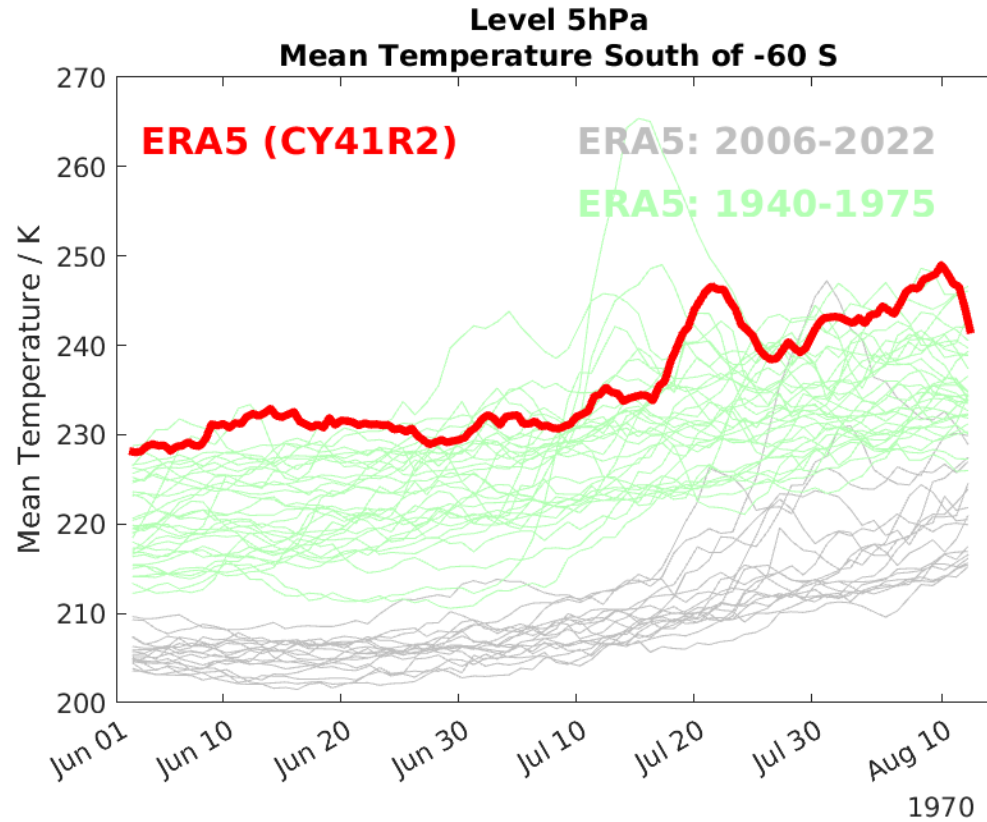


Impact of assimilating IRIS on S. polar stratospheric biases

Climate Change

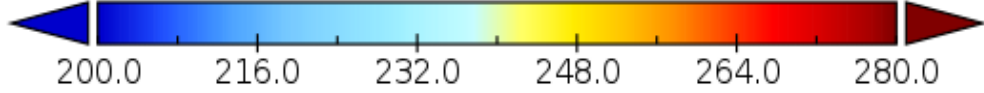


5 hPa temperature
10th July 1970, 00Z



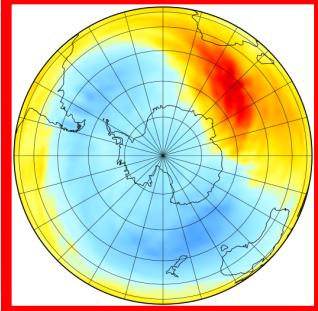
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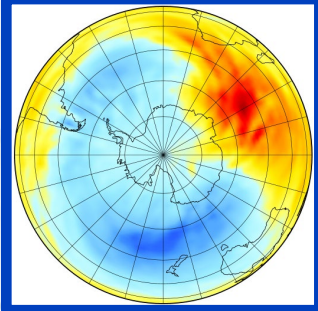


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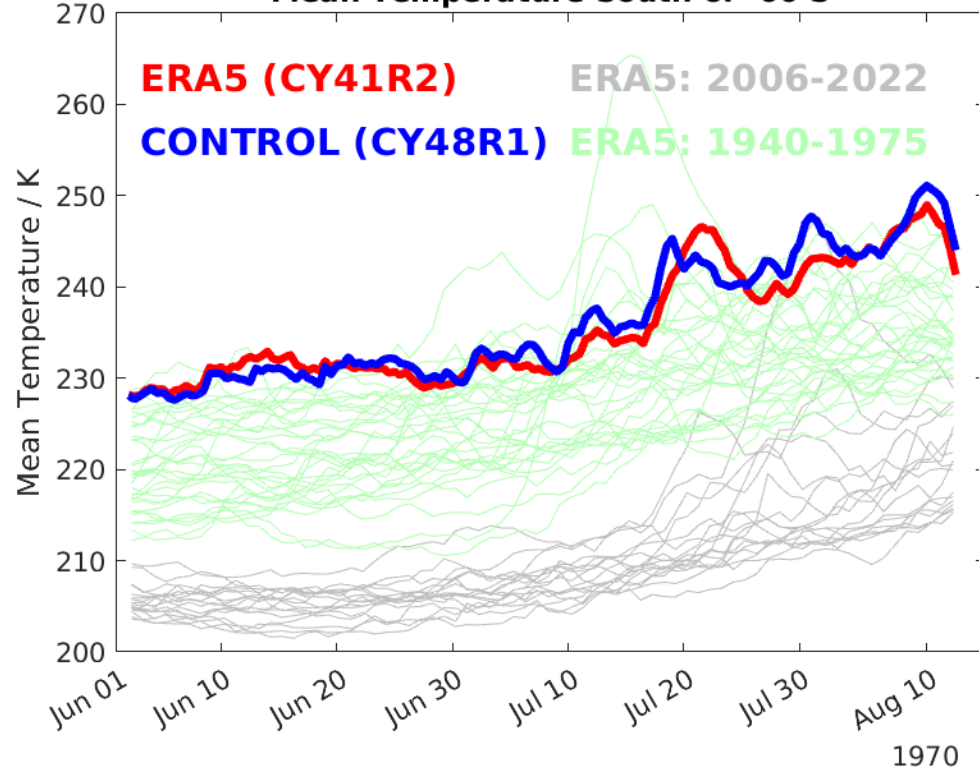
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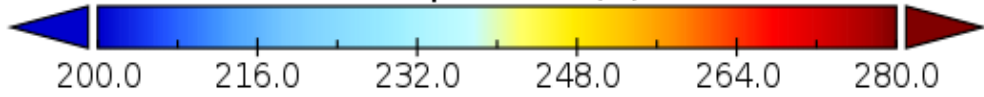


Level 5hPa
Mean Temperature South of -60 S



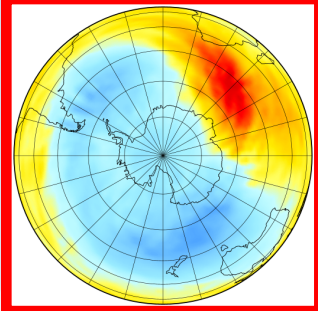
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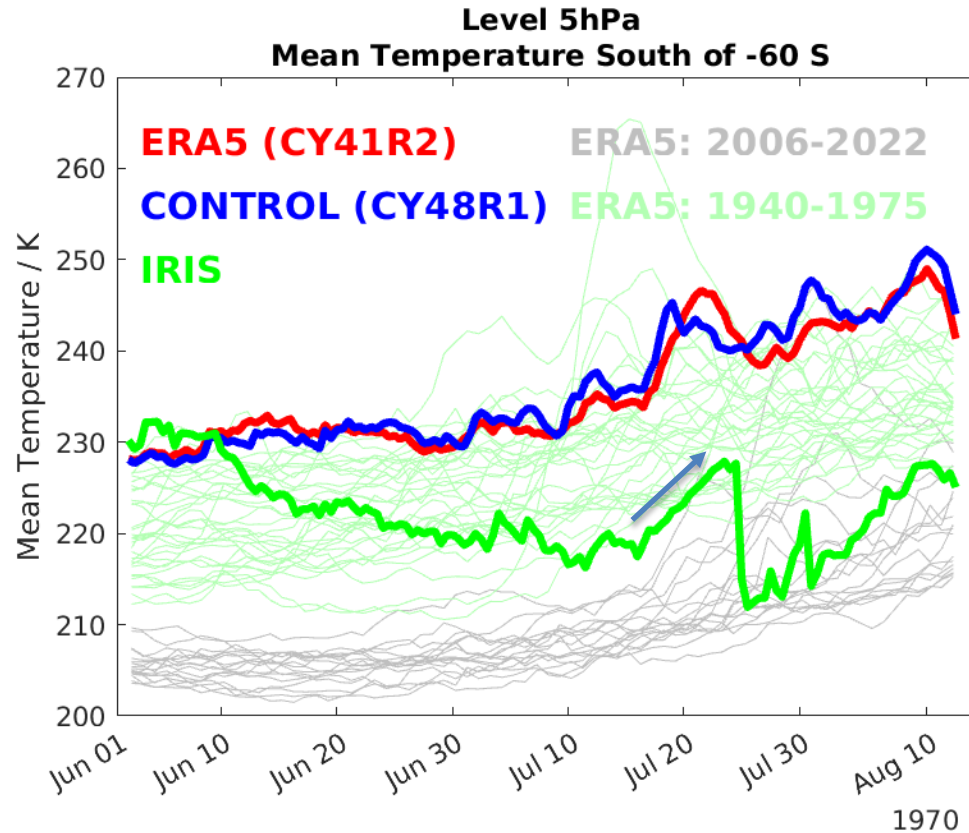
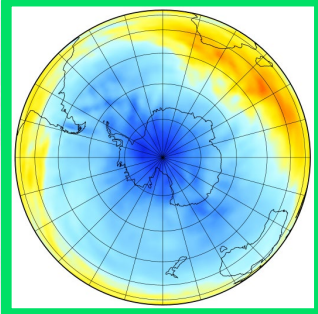
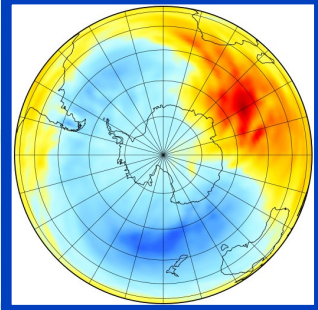


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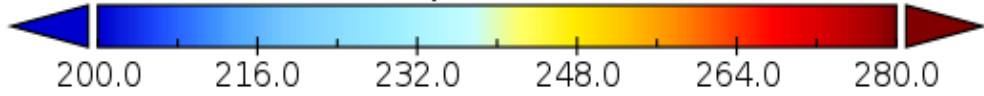


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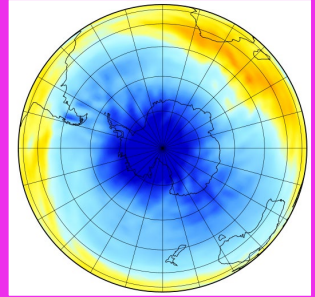
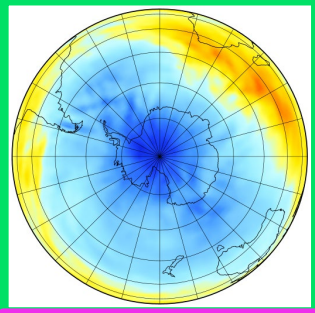
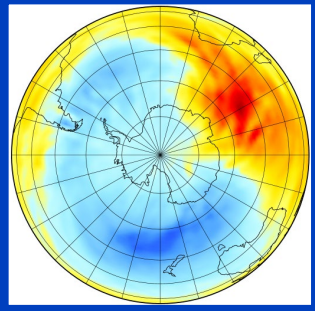
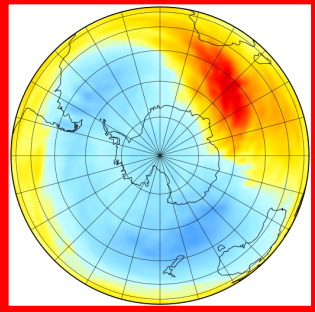
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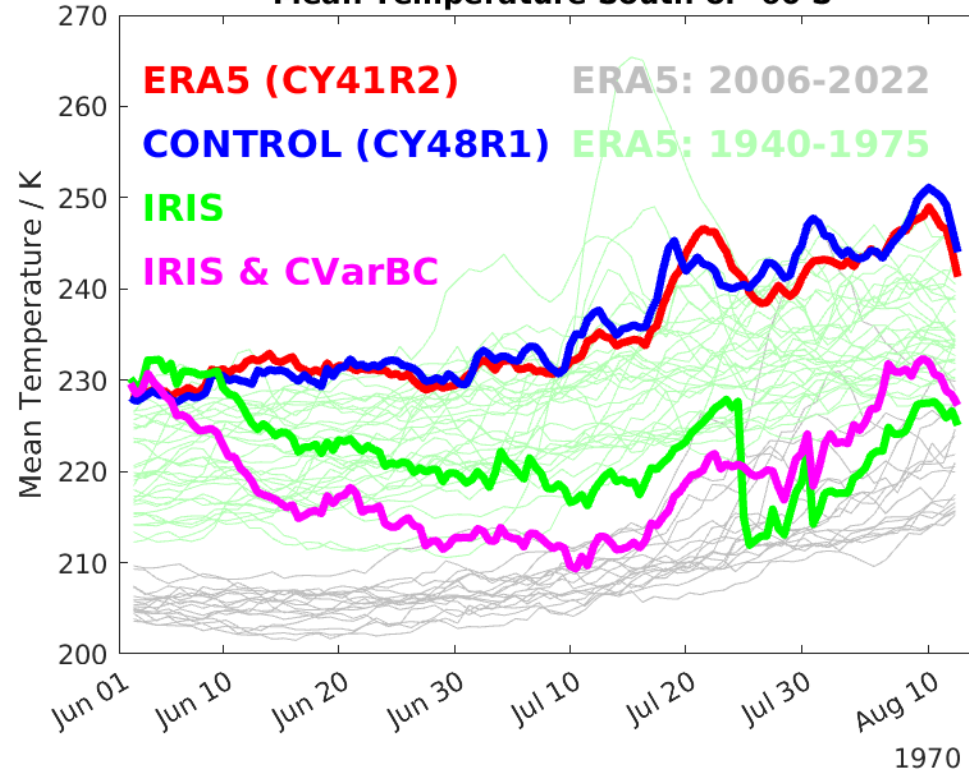
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- **Assimilating IRIS gradually** brings temperatures to more realistic values. Note: increase \nearrow from 16th-24th July is associated with an **outage** of IRIS observations
- Using **Constrained VarBC** (Han & Bormann) reduces the bias absorbed by VarBC, and accelerates cooling of the analysis towards more realistic values.



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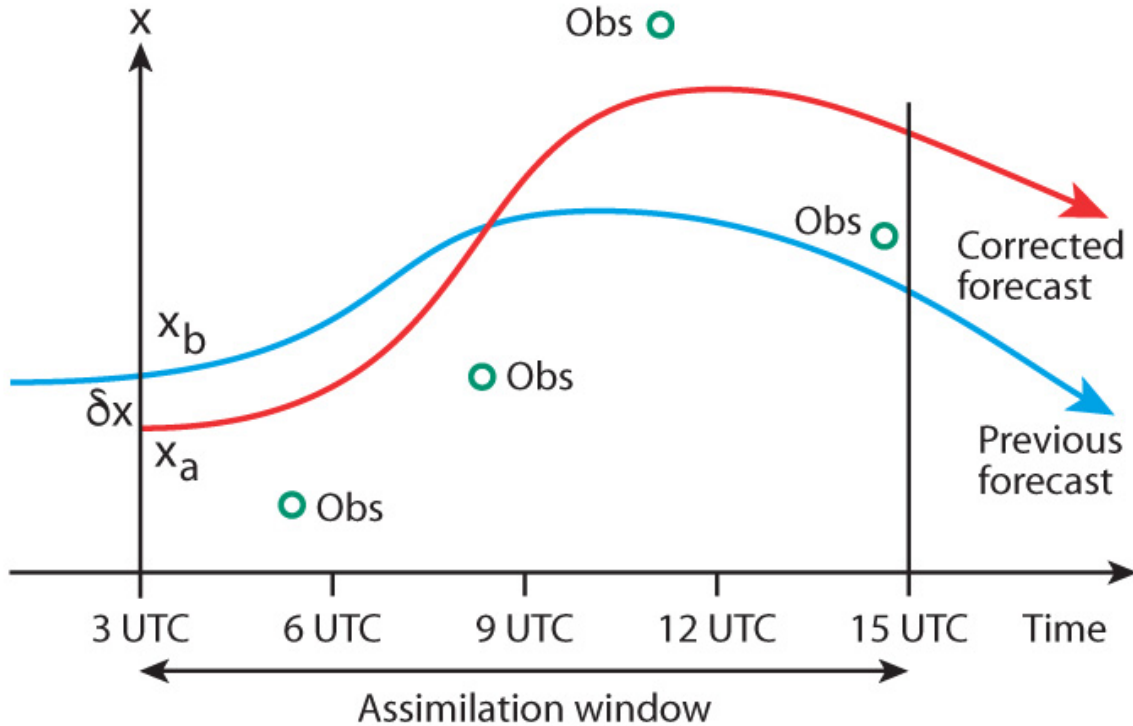
- Weak constraint 4D-Var & model error forcing
- Using early sounding data (IRIS in 1970) to evaluate model error correction strategies

Summary / Future Perspectives



Standard 4D-Var formulation

4D-Var is a common algorithm to find the **optimal initial state** by minimising the discrepancies with the **prior estimate** and the **observations**



Model's equation

$$x_k = \mathcal{M}_k(x_{k-1})$$

4D-Var cost function

$$J(x_0) = \frac{1}{2}(x_0 - x_b)^T \mathbf{B}^{-1}(x_0 - x_b) + \frac{1}{2} \sum_{k=0}^K [y_k - \mathcal{H}(x_k)]^T \mathbf{R}_k^{-1} [y_k - \mathcal{H}(x_k)]$$

- Standard formulation assumes that the model is perfect
- A model trajectory is entirely determined by its initial condition

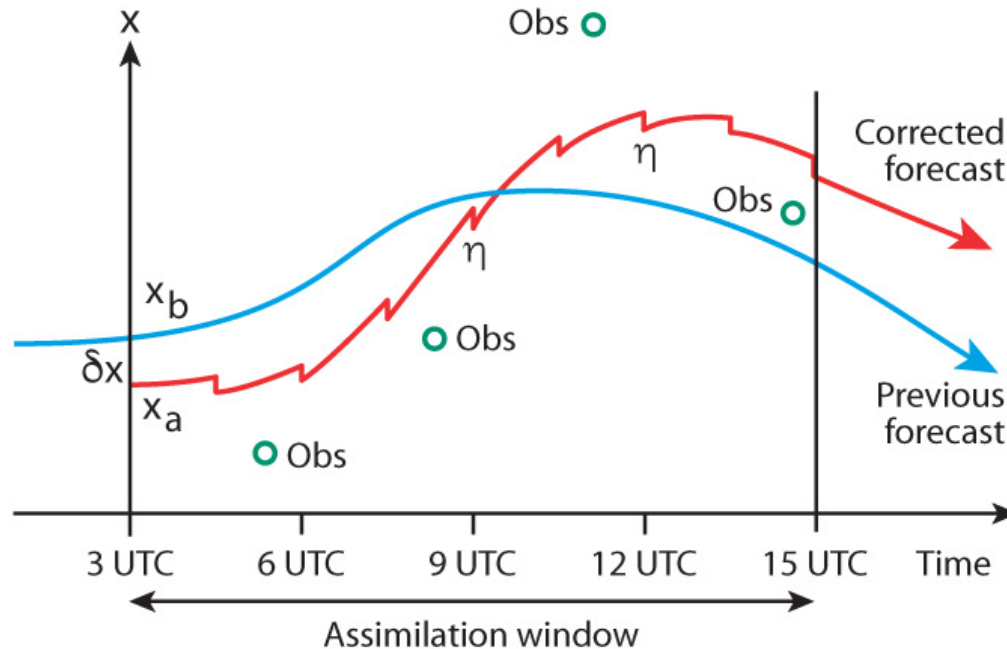


Weak-constraint 4D-Var formulation

We assume that the model is not perfect, adding an error term η in the model equation

$$x_k = \mathcal{M}_k(x_{k-1}) + \eta \quad \text{for } k = 1, 2, \dots, K$$

The model error estimate η contains 3 physical fields (temperature, vorticity and divergence)



- Introduce additional degrees of freedom to fit background and observations
- A model trajectory is entirely determined by its initial condition and the model error forcing
- Concept of scale separation introduced between background and model errors
- Constant model error forcing over the assimilation window



Weak-constraint 4D-Var formulation

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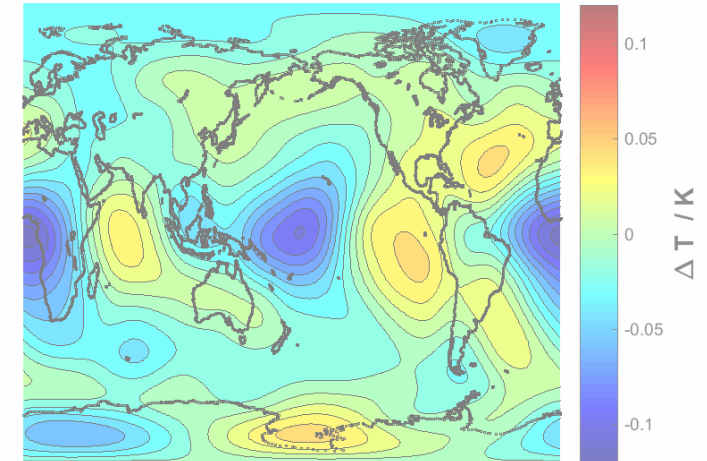
The model error estimate η contains 3 physical fields (temperature, vorticity and divergence)

$$\begin{aligned}
 J(x_0, \eta) &= \frac{1}{2}(x_0 - x_b)^T \mathbf{B}^{-1}(x_0 - x_b) \\
 &+ \frac{1}{2} \sum_{k=0}^K [y_k - \mathcal{H}(x_k)]^T \mathbf{R}_k^{-1} [y_k - \mathcal{H}(x_k)] \\
 &+ \frac{1}{2}(\eta - \eta_b)^T \mathbf{Q}^{-1}(\eta - \eta_b)
 \end{aligned}$$

Model initial condition \uparrow

Model bias correction \uparrow

Weak Constraint 4DVar Model Error Estimate at 5hPa
01-Dec-2019 00:00:00

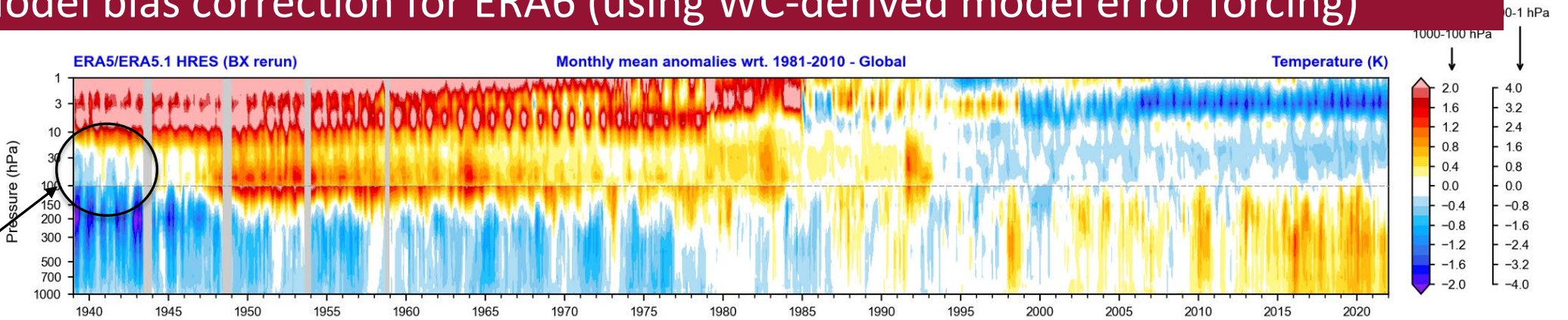


- ➔ Introduce additional degrees of freedom to fit background and observations
- ➔ A model trajectory is entirely determined by its initial condition and the model error forcing
- ➔ Concept of scale separation introduced between background and model errors
- ➔ Constant model error forcing over the assimilation window



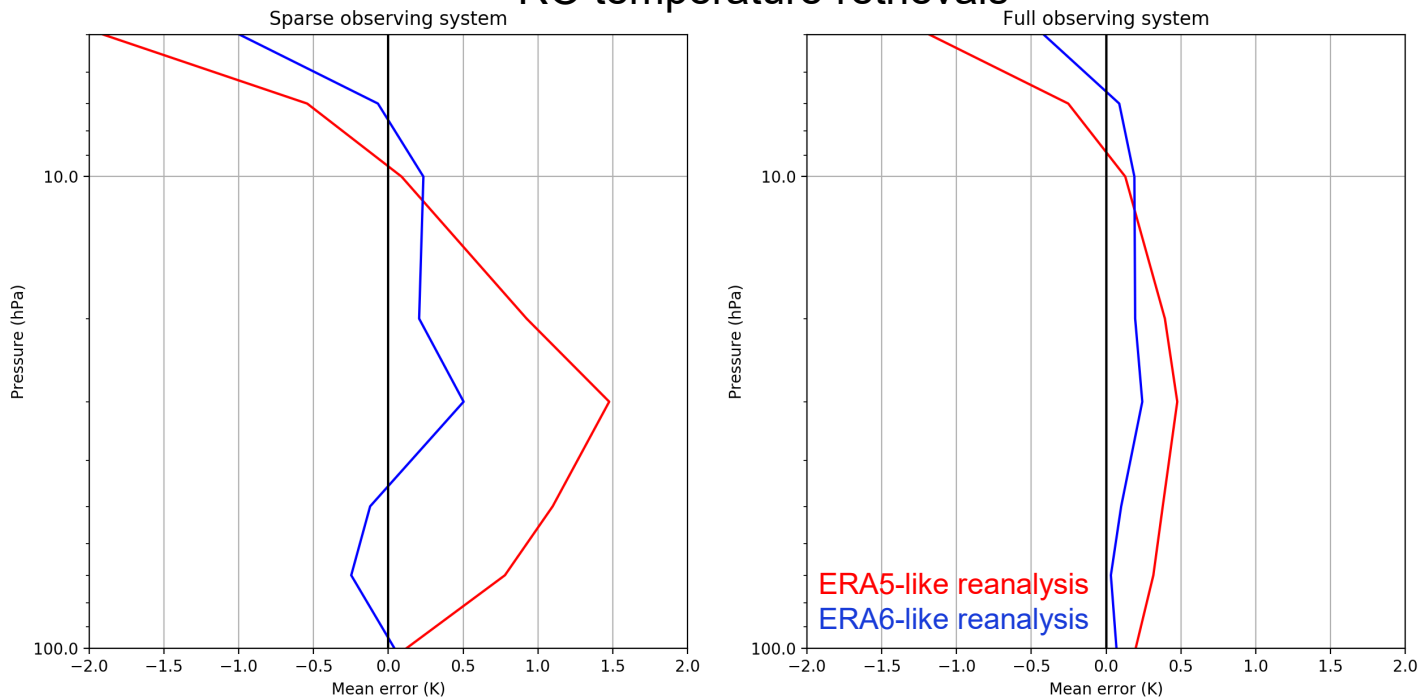
Climate Change

Model bias correction for ERA6 (using WC-derived model error forcing)



Extension of ERA5 to 1940
 Few upper air observations exposed the model bias

First guess mean error with respect to RO temperature retrievals

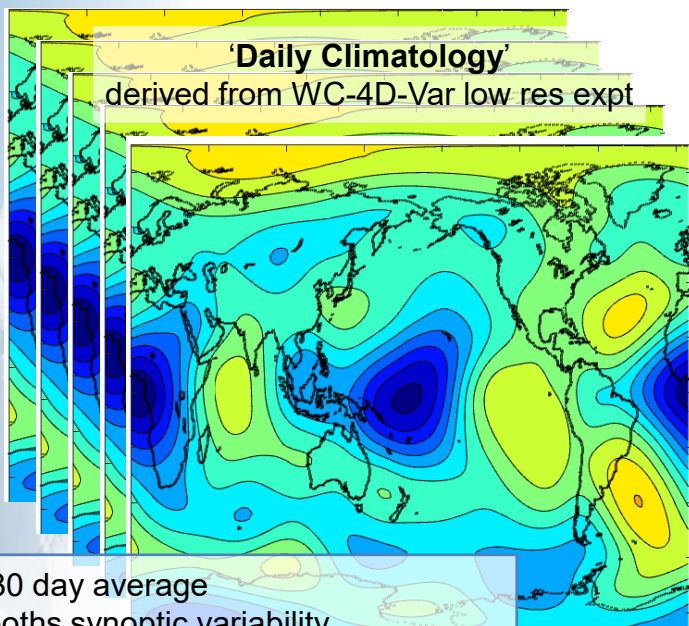


Recent observing system

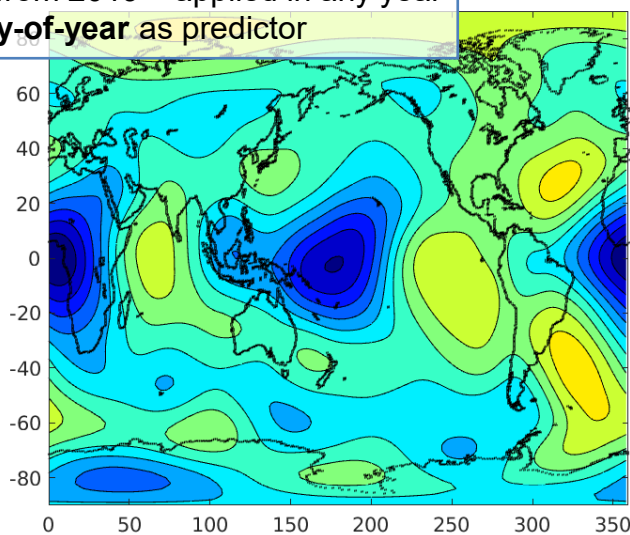


Climate Change

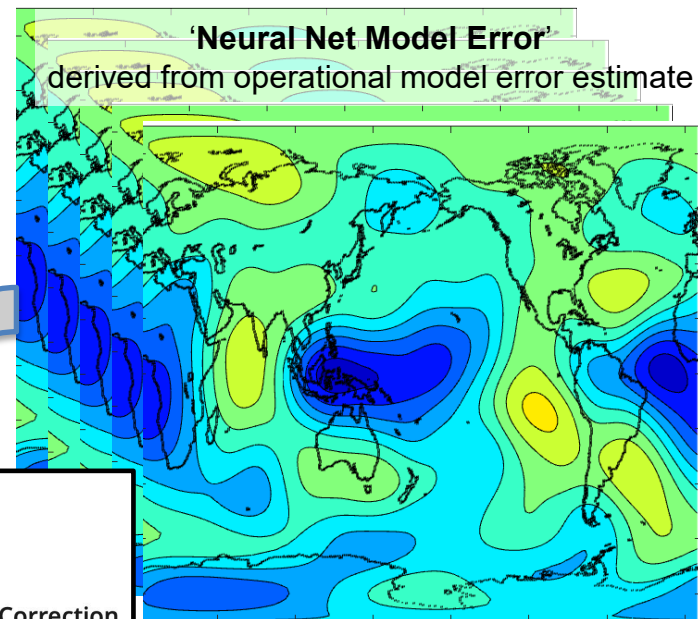
Model error climatology derived from weak constraint 4D-Var estimates of model error



- +/- 30 day average
- smooths synoptic variability
- TCo399 resolution (28km)
- derived from 2019 – applied in any year
- uses **day-of-year** as predictor



model error estimates at 5hPa



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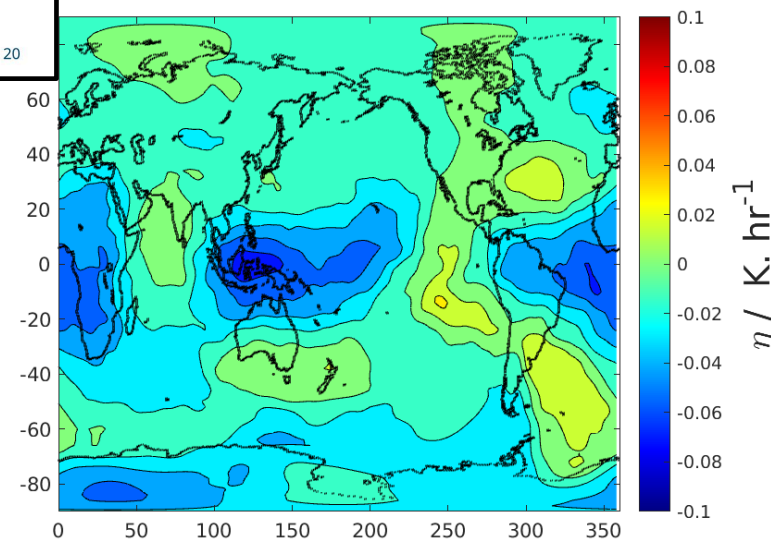
Machine Learning for Model Error Inference and Correction

Massimo Bonavita ✉ Patrick Laloyaux

First published: 13 November 2020 | <https://doi.org/10.1029/2020MS002232> | Citations: 20

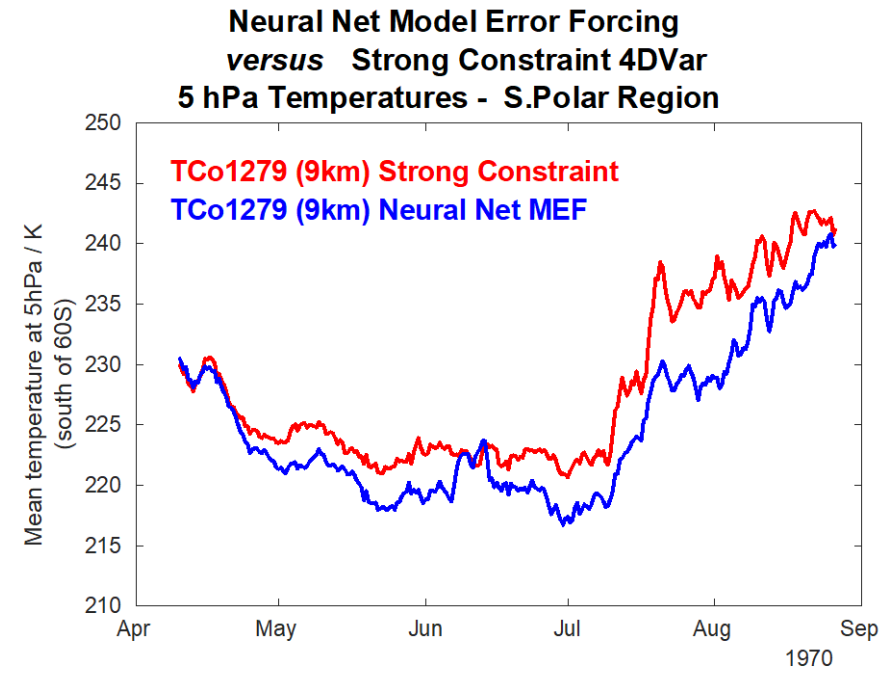
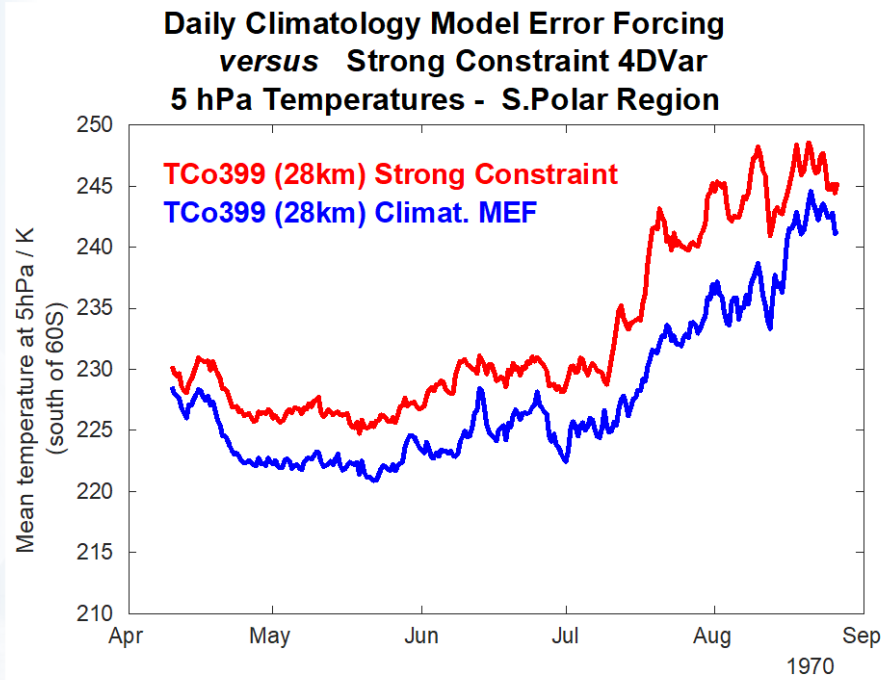
- derived from June 2020 - Feb2022
- predictors:
 - lat & lon
 - time of day & month of year
 - background state (T)

estimates shown for 1/7/1970





Model error forcing experiments in 1970 – impact on upper stratospheric temperatures

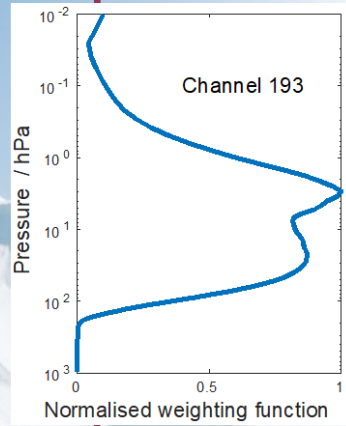


- For strong constraint & model error forcing experiments: increase in resolution (28km to 9km) helps lower minimum temperatures (230K->223K in June 1970)
- Model error forcing (both types) results in additional cooling of ~5K, with minimum temperatures of 217K
- ... but doesn't bring temperatures to the minimum temperatures expected (from IRIS assimilation experiments) of ~210K
- expect ERA6 (TCO799) will be closer to behavior of TCo1279 experiment shown here.

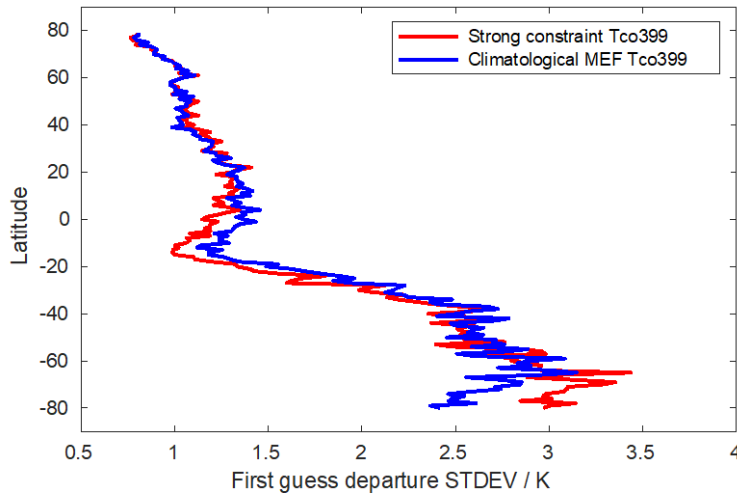
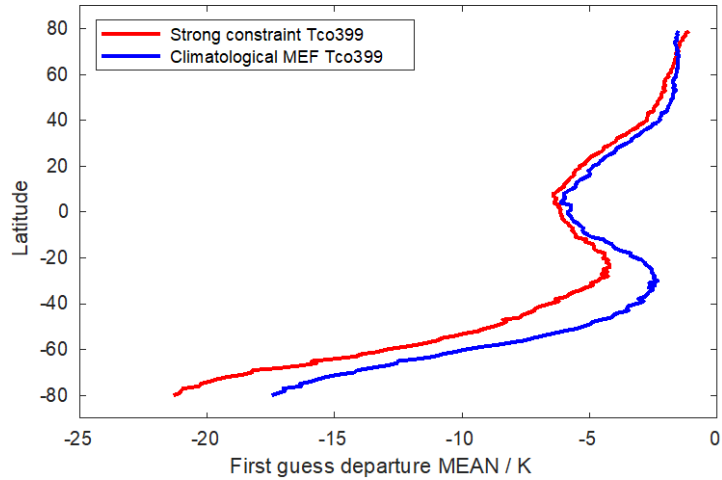


Climate Change

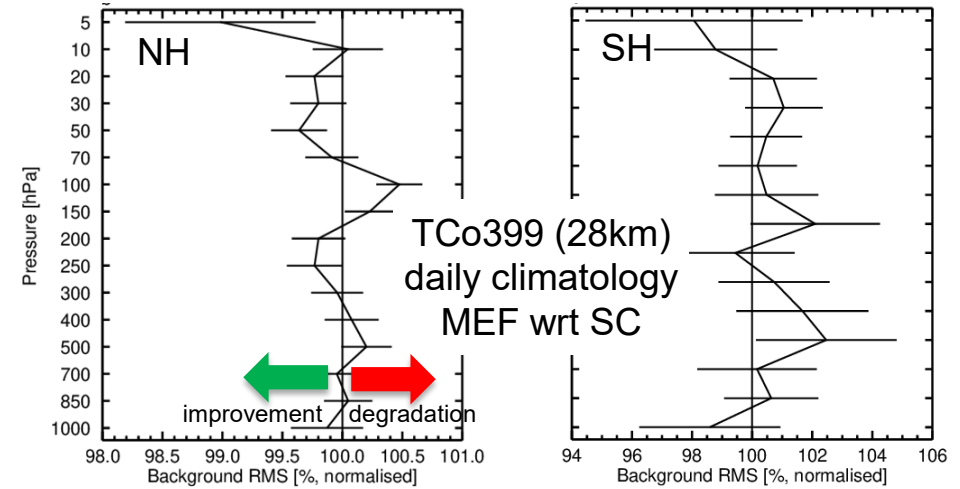
Verification of impacts of MEF: background fits to IRIS and radiosondes



Impact of model error forcing versus strong constraint 4DVar on (passive) IRIS channel 193 fg departures July 1970



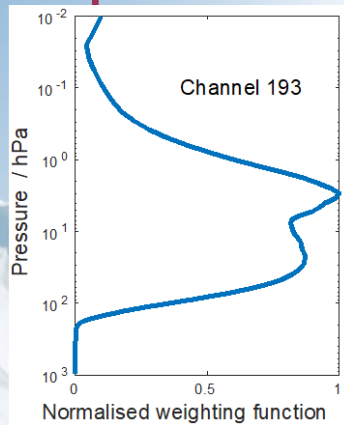
Background fits to radiosonde temperatures 20th April – 26th August 1970



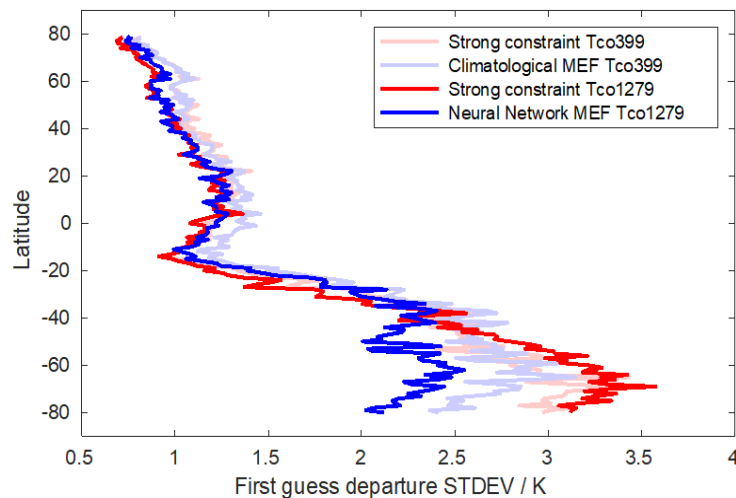
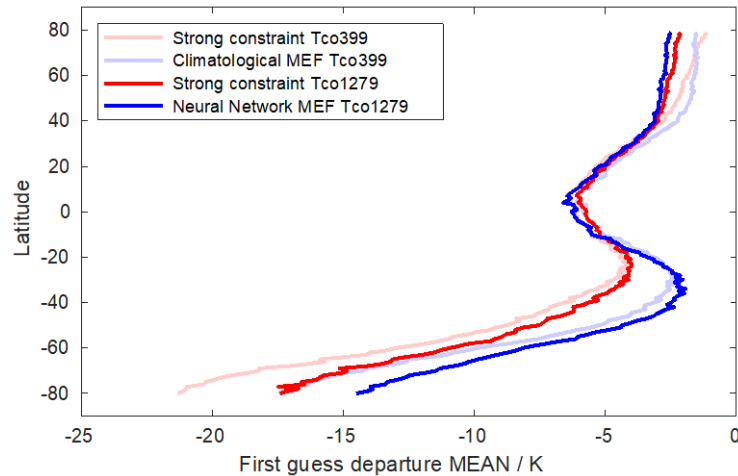


Climate Change

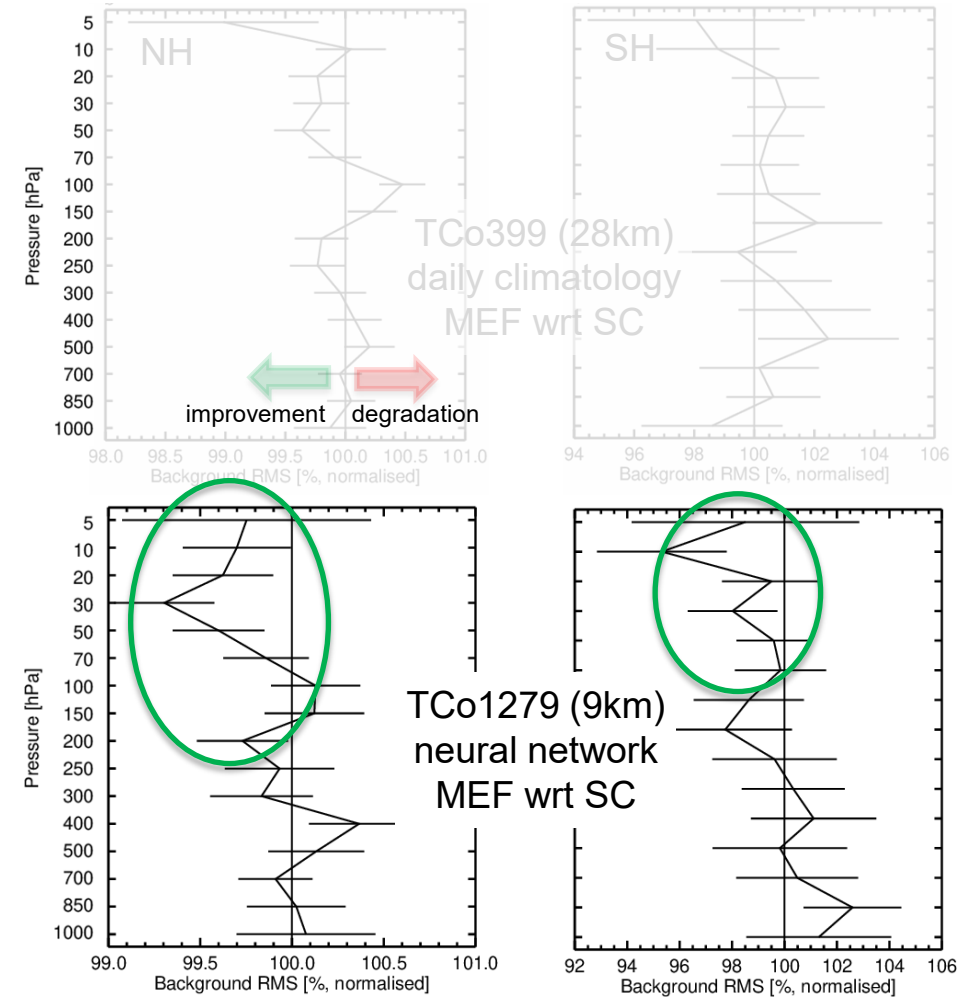
Verification of impacts of MEF: background fits to IRIS and radiosondes



Impact of model error forcing
versus strong constraint 4DVar
on (passive) IRIS channel 193 fg departures
July 1970



Background fits to radiosonde temperatures
20th April – 26th August 1970



- NN MEF improves bias and synoptic performance
- IRIS provides unique insight into biases in otherwise observation sparse domains
- But significant biases remain



Biases in stratospheric temperatures in ERA5

- In '*model space*': anomalies & analysis increments
- ERA 5.1 & the role of RO observations
- In '*observation space*': mean first guess departures in ERA5, ERA-Interim & *proto-ERA6*

Southern winter polar bias in ERA5

- Exposed by anomalies & IRIS data

Exploiting information from RO observations back in time

- Weak constraint 4D-Var & model error forcing
- Using early sounding data (IRIS in 1970) to evaluate model error correction strategies

Summary / Future Perspectives



Summary & conclusions

- Biases in stratospheric temperatures are particularly evident in ERA5. GNSSRO data:
 - **Has played a key role in mitigating the effects of these biases** in the recent (2006 →) era, and ;
 - **Will play a role in mitigating their effects in earlier epochs of ERA6** (1950 → 2006) through WC-4DVar & model error forcing
- The magnitude of the biases is large (typically ~1K, but up to 20K!). In successive generations of reanalyses, attention will turn to much smaller biases in other regions (& variables). We hope that the diagnostics and methods used to mitigate in ERA5 & ERA6 will be useful in those cases
- Short lived early satellite missions (*e.g.* IRIS, in 1970) have proved valuable in assessing the performance of model error forcing, by providing observations in otherwise unobserved regions/domains
- ERA6 will make use of reprocessed RO datasets for COSMIC, CHAMP, GRACE and GRAS – provided by EUMETSAT. Impacts (not shown here) are generally positive
- RO data , and other reference datasets (*e.g.* GRUAN radiosondes & CrIS radiances), perhaps have a role to play in evaluating uncertainties in ERA6 (withhold a subset of RO observations, and use to validate the reanalysis ?)



Climate
Change



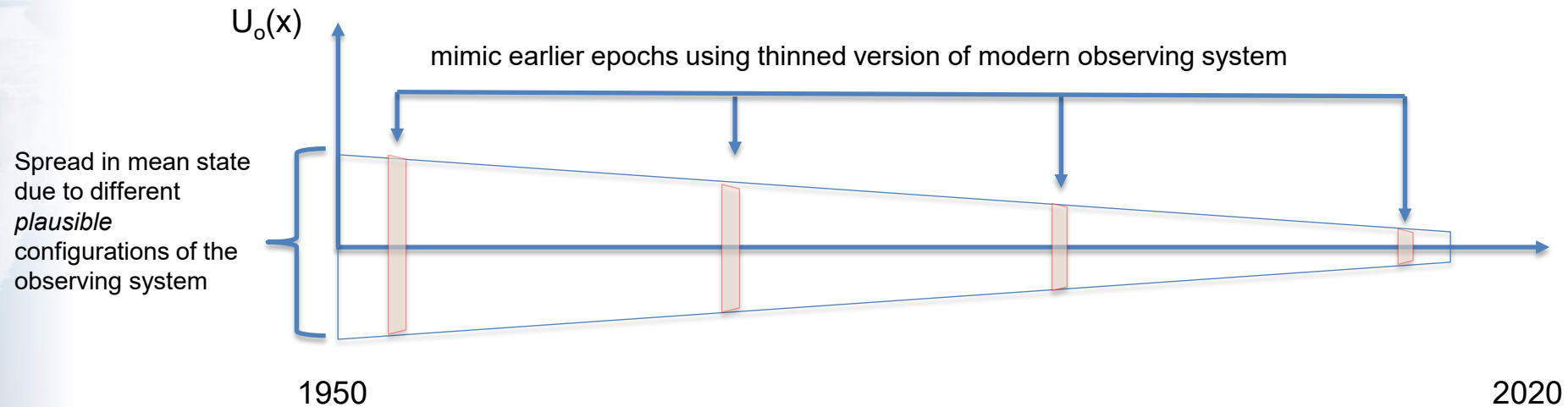
Extra slides



Possible approaches to determining mean-state uncertainty

The observing system component

- Defined here as - “*uncertainty in mean state arising from uncorrected biases in the observing system & choice of observing system configuration*”
- OSEs with different plausible configurations of observing system, for each epoch
- Simplest approach: withdraw ‘redundant’ components of observing system and evaluate change in the mean state (next slide)
- Other factors: choice of observational data, bias model, QC/thinning, observation errors, ...

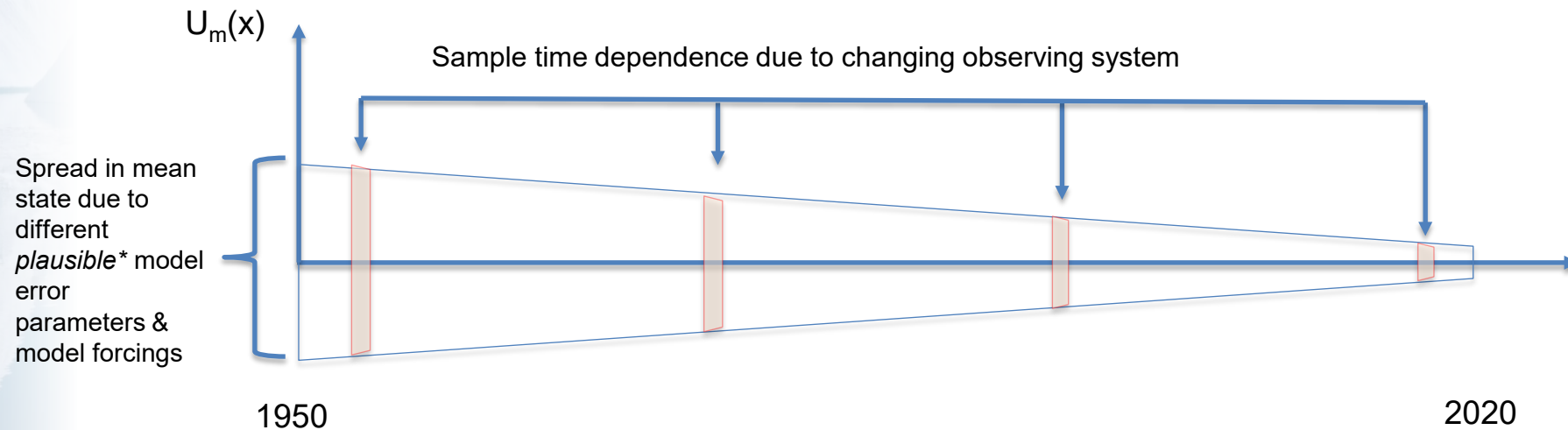




Possible approaches to determining mean-state uncertainty

The model component

- Defined here as - “*uncertainty in mean state arising from uncertain model parameters and forcings*”
- Changes in time, due to the changing observing system
- OSEs with perturbed model parameters & alternative choices of forcings
- Key model parameters? - draw upon experience of EPS and climate modelling communities
- Sample time dependence using paired down modern observing system, or run in past epochs

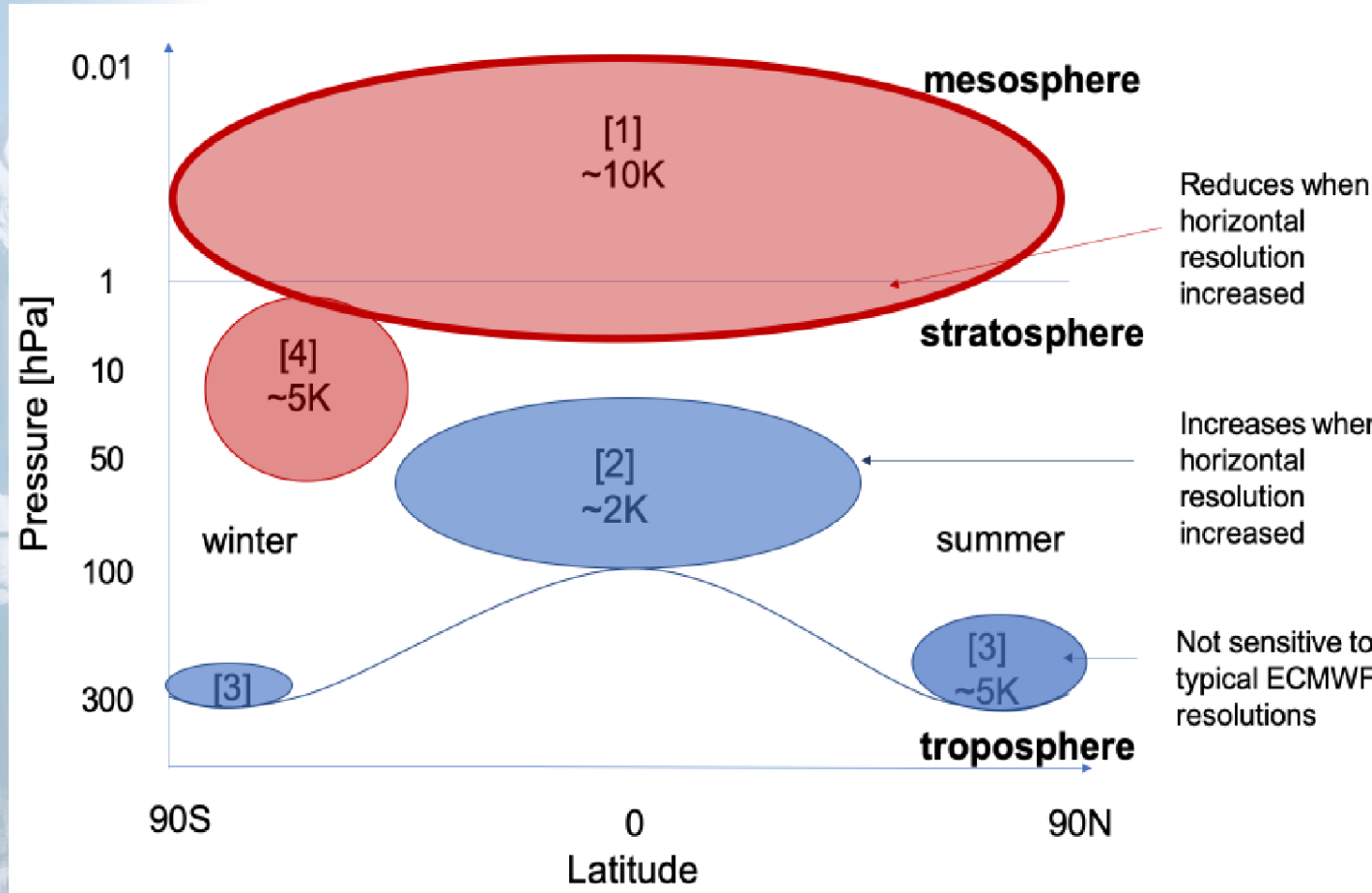


- Perturbed by magnitudes consistent with documented uncertainties and/or giving rise to no significant degradation in forecast skill in OSEs



Climate
Change

Model biases in the stratosphere



Weak constraint 4D-Var offers a solution for ERA6.

In addition, future improvements are expected from :

[1] – revised radiation scheme, improved SW solar spectrum, improved (and interactive) ozone,

[2] improved dynamical core

[3] reduction of H₂O in lower stratosphere, improved methane oxidation scheme